Interpretation Report 8: December 2016 to May 2017

Scarborough
Borough Council

October 2017





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Contents

Section		Page
Disclaimer		viii
Summary of	findings	ix
Introduction		1-1
1.1	Background to study	1-1
1.2	Aims and objectives of monitoring	1-1
1.3	Programme of work	1-2
1.4	Scope of data analysis work	
1.5	Cliff instability hazard assessment	
1.6	Checks of monitoring equipment integrity	1-4
Weather Sur	mmary	2-1
2.1	Introduction	2-1
	2.1.1 Rainfall and landslides	2-5
2.2	Summary	2-5
Runswick Ba	ıy	3-7
3.1	Site description	
3.2	Ground model and monitoring regime	
3.3	Historical ground behaviour	
3.4	New data	3-8
3.5	Causal response relationships	3-8
3.6	Implications and recommendations	3-8
Whitby West	t Cliff	4-1
4.1	Site description	
4.2	Ground model and monitoring regime	
4.3	Historical ground behaviour	
4.4	New data	
4.5	Causal-response relationships	4-2
4.6	Implications and recommendations	
Robin Hood'	s Bay	5-1
5.1	Site description	
5.2	Monitoring regime	
5.3	Historical ground behaviour	
5.4	New data	
5.5	Causal-response relationships	
5.6	Implications and recommendations	
Scalby Ness.		6-1
6.1	Site description	
6.2	Ground model and monitoring regime	
6.3	Historical ground behaviour	
6.4	New data	
6.5	Causal-response relationships	
6.6	Implications and recommendations	
Scarborough	n North Bay – Oasis Café	7.0

7.1	Site description	7-9
7.2	Ground model and monitoring regime	7-9
7.3	Historical ground behaviour	7-9
7.4	New data	7-9
7.5	Causal-response relationships	7-10
7.6	Implications and recommendations	7-10
Scarborough	North Bay – The Holms	8-1
8.1	Site description	
8.2	Ground model and monitoring regime	
8.3	Historical ground behaviour	8-1
8.4	New data	8-2
8.5	Causal-response relationships	8-3
8.6	Implications and recommendations	8-3
•	South Bay	
9.1	Site description	
9.2	Ground model and monitoring regime	
9.3	Historical ground behaviour	
9.4	New data	
	9.4.1 St Nicholas Cliff (MU 22A)	
	9.4.2 Spa Chalet (MU 22/1)	
	9.4.3 Spa (MU 22/2 and 22/3)	
	9.4.4 Clock Café (MU 22/4)	
	9.4.5 South Cliff Gardens (MU 22/5 and 22/6)	9-20
	9.4.6 Holbeck Gardens (MU 22/7)	9-24
9.5	Causal-response relationships	9-25
9.6	Implications and recommendations	9-26
Filey Town		10-27
10.1	Site description	10-27
10.2	Ground model and monitoring regime	10-27
10.3	Historical ground behaviour	10-28
10.4	New data	10-28
10.5	Causal-response relationships	10-33
10.6	Implications and recommendations	
Filey Flat Cliff	S	11-1
11.1	Site description	11-1
11.2	Ground model and monitoring regime	11-1
11.3	Historical ground behaviour	11-1
11.4	New data	11-2
11.5	Causal-response relationships	11-5
11.6	Implications and recommendations	11-6
References		12-1

Appendix

A Monitoring data (DVD format only)

Tables

1.1	Monitoring	1	
1.1	שווו ווווווווווווווו	เบเสเเบาเร	מווט טבעונבז

- 1.2 Programme of data collection and reporting
- 1.3 Instability hazard assessment guidance level
- 2.1 Monthly rainfall recorded at Flat Cliffs and Scarborough Spa meteorological stations
- 3.1 Summary of historical ground behaviour at Runswick Bay
- 3.2 Summary of inclinometer data at Runswick Bay
- 4.1 Summary of historical ground behaviour at Whitby West Cliff
- 4.2 Summary of inclinometer data at Whitby West Cliff
- 5.1 Summary of historical ground behaviour at Robin Hood's Bay
- 5.2 Summary of inclinometer data at Robin Hood's Bay
- 5.3 Summary of groundwater data at Robin Hood's Bay
- 6.1 Summary of historical ground behaviour at Scalby Ness
- 6.2 Summary of inclinometer data at Scalby Ness
- 6.3 Summary of groundwater data at Scalby Ness
- 7.1 Summary of historical ground behaviour at Oasis Café
- 7.2 Summary of inclinometer data at Oasis Café
- 7.3 Summary of groundwater data at Oasis Café
- 8.1 Summary of historical ground behaviour at The Holms
- 8.2 Summary of inclinometer data at The Holms
- 8.3 Summary of groundwater data at The Holms
- 9.1 Summary of historical ground behaviour at Scarborough South Bay
- 9.2 Summary of inclinometer data at St Nicholas Cliff
- 9.3 Summary of groundwater data at St Nicholas Cliff
- 9.4 Summary of inclinometer data at Spa Chalet
- 9.5 Summary of groundwater data at Spa Chalet
- 9.6 Summary of inclinometer data at the Spa
- 9.7 Summary of groundwater data at the Spa
- 9.8 Summary of inclinometer data at the Clock Café
- 9.9 Summary of groundwater data at the Clock Café
- 9.10 Summary of inclinometer data at South Bay Gardens
- 9.11 Summary of groundwater data at South Bay Gardens
- 9.12 Summary of inclinometer data at Holbeck Gardens
- 9.13 Summary of groundwater data at Holbeck Gardens
- 10.1 Summary of historical ground behaviour at Filey Town
- 10.2 Summary of inclinometer data at Filey Town
- 10.3 Summary of groundwater data at Filey Town
- 11.1 Summary of historical ground behaviour at Flat Cliffs
- 11.2 Summary of inclinometer data at Flat Cliffs
- 11.3 Summary of groundwater data at Flat Cliffs

Figures

- 2.1 Comparison of monthly rainfall records (2011 to 2017).
- 2.2 Daily rainfall recorded at Flat Cliffs during 2017
- 2.3 Seasonal rainfall comparison (2011-2017)
- 2.4 Maximum daily wind speed (2011 to 2017)
- 2.5 Air temperature variation (2011 to 2017)
- 2.6 Monthly rainfall and antecedent totals (2011 to 2017). Ground movements were recorded at Scalby Mills during December 2012 (red box).
- 3.1 Location of slope monitoring at Runswick Bay
- 4.1 Location of slope monitoring at Whitby West Cliff

- 5.1 Location of slope monitoring at Robin Hood's Bay
- 6.1 Location of slope monitoring at Scalby Ness
- 7.1 Location of slope monitoring at Scarborough North Bay Oasis Café
- 8.1 Location of slope monitoring at Scarborough North Bay The Holms
- 9.1 Location of slope monitoring at Scarborough South Bay
- 9.2 Cumulative AE (RDC) and AE rate (RDC/hr) time series at Scarborough Spa for the period December 2016 to September 2017.
- 9.3 Cumulative AE (RDC) and cumulative rainfall time series at Scarborough Spa for the period December 2016 to September 2017.
- 10.1 Location of slope monitoring at Filey
- 11.1 Location of slope monitoring at Flat Cliffs
- 11.2 Cumulative AE (RDC) and AE rate (RDC/hr) time series at Flat Cliffs, Filey for the period December 2016 to September 2017.
- 11.3 Cumulative AE (RDC) and cumulative rainfall time series at Flat Cliffs, Filey for the period December 2016 to September 2017.

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The interpretation of the level of cliff instability risk presented in this document is based solely on the data provided by JBA. While every effort will be made to ensure the data are correct, Halcrow cannot be held responsible for the quality of monitoring data. This data analysis report comments on the monitoring data collected over the preceding 6 month period at specific locations. It will not make projections of future cliff instability activity or discuss cliff instability risk at areas that are not monitored. It is Scarborough Borough Council's responsibility to determine an appropriate response to the guidance on cliff instability risk provided in this report.

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Summary of findings

This report presents an interpretation of coastal slope monitoring data recorded between December 2016 and May 2017 along the Scarborough Borough Council frontage. It is the eighth in a series of 6-monthly updates on the cliff instability risk of the frontage that began in 2014. The weather over the winter 2016/2017 and spring 2017 period covered by this report was generally drier than average. Boreholes show that water levels have remained at previous low levels or fallen slightly during the monitoring period, except for Scalby Ness (B9), Oasis Café (BH4p), Scarborough Spa (BH104b), South Cliff Gardens (BH18a) and Filey Town (CPBH01a, CPBH02a and CPBH10a) that are at atypically high levels. In situ monitoring using inclinometers does not indicate any slope movement, indicating that localised elevated groundwater levels have not triggered ground movement. Experimental in situ monitoring using Acoustic Emissions devices installed by Loughborough University at Scarborough Spa and Flat Cliffs also show negligible slope movement.

Specific sites of concern and issues needing attention are as follows:

- At Robin Hood's Bay, BH1a shows groundwater levels have remained steady at lower levels that were seen prior to the installations being inadvertently covered with tarmac. It is likely that this borehole had been affected by surface water ingress, but now appears to be functioning correctly. Boreholes BH1b and BH3b were dry, and the integrity of the piezometers should be checked. Data is unavailable for inclinometer BH2 for this monitoring period as the borehole was obstructed by a vehicle at the time of the site readings.
- At Scalby Ness, the groundwater level in borehole Sn2b has fallen significantly from the historical high recorded in the previous monitoring period, following a dry winter and spring. This site should continue to be visually monitored, particularly if rainfall is high during the summer and autumn. Borehole B9 experienced a rise in groundwater level and this trend should also be reviewed in the next monitoring report. Piezometers P1a, P1b and P3 experienced a problem downloading data and should be investigated and remedied by the monitoring contractor at the next download. Piezometers P4a and P4b record changes in groundwater with the same pattern but at slightly different levels, which suggests piezometers are not monitoring discrete horizons and the seal between the two piezometers is damaged. Boreholes B6 and Sn2a were dry, suggesting the piezometers are damaged. These locations should continue to be monitored.
- At Oasis Café, attention should be given to piezometer BH4p where elevated groundwater during drier than average months suggest an ingress of surface water from above the cliff or a local source of groundwater not associated with rainfall recharge, such as leaking services.
- At Scarborough North Bay (The Holms), data collection has resumed in BH9a following problems downloading the data in the previous monitoring periods. Groundwater levels were well below the historical high and remain steady. Piezometer BH9b is no longer functional and should be repaired.
- At Scarborough Spa Chalet, no data has been recorded beyond the two-previous monitoring period at piezometer BH12. This site requires attention to fix or replace the piezometer and damaged cable.
- At Scarborough Spa, groundwater levels have remained steady or are falling, except at BH104b where water level remains at an historical high. No movement was detected at inclinometers nearby, but this site should be visually monitored, particularly following sustained wet weather. Several boreholes were dry (5 Spa, G1a, G1b, BH106a, BH106b and BH108b), suggesting the piezometer installations are damaged. These locations require attention and should continue to be monitored. Data loggers installed at BH1 (Prom) show a potential systematic error in the data and should be checked.

- At the Clock Café, borehole BH15 remains dry, and the integrity of the piezometer should be checked. Inclinometer AA10 readings were not available during this monitoring report and should be collected on the next site visit.
- At South Cliff Gardens, shallow piezometer BH18a shows a rapid increase in groundwater levels relative to deeper piezometer BH18b which shows a more gradual increase. This suggests surface water ingress and the contractor should ensure that plastic caps are in place and that water cannot collect at the top of the boreholes. No movement has been indicated by the inclinometers in this location. At borehole D2a, D2b, BH3a and BH3b there was a problem downloading data. The issue should be investigated and remedied by the monitoring contractor next time.
- At Holbeck Gardens there was a problem downloading data at borehole BH4a and BH4b. The issue should be investigated and remedied by the monitoring contractor next time.
- At Filey Town, in boreholes CPBH02a and CPBH09a groundwater levels remain near historical highs, levels atypical given the dry conditions during this monitoring period. At CPBH01a groundwater levels have risen significantly again following a fall during the previous monitoring period, which may suggest ingress of surface water. There is also a rise in groundwater level in CPBH10a towards the historical high. These locations should be visually checked and reviewed in the next monitoring period, particularly if sustained wet weather occurs. Boreholes CPBH02b, CPBH08b and CPBH10b are dry suggesting the piezometer installation may be damaged, however, the location should continue to be monitored. There was a problem downloading data from CPBH01b, CPBH04, CPBH06a and CPBH06b. This issue should be investigated and remedied by the monitoring contractor during the next visit. Borehole CPBH03 was obstructed during the site visit and data should be collected for the next monitoring report.
- At Filey Flat Cliffs there was a problem downloading data at borehole C4a. The issue should be investigated and remedied by the monitoring contractor next time.

Introduction

1.1 Background to study

The Scarborough Borough Council coastline is affected by widespread cliff instability, largely due to its geology and climate. Since the Holbeck Hall landslide in June 1993, understanding the risk posed by landslides has been a high priority for the Council. Numerous ground investigations and associated studies at locations of particular concern have been undertaken in the last 20 years meaning the Council now has a widespread network of ground monitoring instrumentation installed, much of which is automated using data-loggers. The Council has also supported the installation of experimental acoustic inclinometers by Loughborough University along its frontage. These experimental devices have the potential to provide cost-effective and accurate real time information on ground movement. The dataset allows the Council to better understand cliff instability risk and support decisions on risk management.

A comprehensive programme of data collection and analysis was commenced by the Council in October 2008, when SBC awarded Mouchel Ltd a contract to design a monitoring strategy for the coastline. Mouchel's recommendations were adopted by SBC and a four-year contract for regular data collection and monitoring reports was awarded. The 7th and final of these reports covered the period up to spring 2012, and was issued in August 2012 (Mouchel 2012).

On completion of this contract, SBC commissioned Haskoning UK Ltd to undertake a thorough review of the condition of boreholes and associated monitoring instruments (Haskoning, 2013). This report highlighted a number of instruments were damaged, due to shearing of the borehole, wear and tear, and vandalism. The work allowed SBC to develop a revised list of instruments and prepare tender documents for re-tendering of data collection and analysis work.

SBC invited tenders on 24 July 2013, with separate contracts for data collection and data analysis being let. Contracts covering an initial three-year programme were awarded on 3 September 2013 to JBA Consulting Ltd and Halcrow Group Ltd (a CH2M company), for data collection and data analysis respectively. JBA undertook the first data collection exercise in November 2013. Data analysis is reported in separate CH2M reports (CH2M 2014a & b, 2015a & b, 2016a). A two-year extension to the project was awarded to the incumbent team in March 2016.

This report provides the seventh set of data analysis and is presented as a stand-alone document.

1.2 Aims and objectives of monitoring

The main objective of the monitoring programme is to provide property- and land-owners with information on instability hazard and risk in vulnerable areas.

The sites and monitoring devices covered by this work are summarised in Table 1.1. Note that some boreholes may have multiple piezometers installed in order to monitor multiple water tables, inclinometers and piezometers are never located in the same boreholes and water-levels are not recorded in boreholes instrumented with inclinometers.

To meet this objective, the specific aims of the study are as follows:

- To place the preceding 6 months' monitoring data in the context of the historical record
- To highlight the implications of the data to coastal instability risk management

In addition, the ultimate aim of the study is:

 To collect sufficient monitoring data to enable site-specific relationships between rainfall, groundwater levels and ground movement to be understood. With sufficient data, it is hoped that threshold rainfall and groundwater levels, above which instability is likely to be triggered, can be identified. This understanding will eventually allow early warning of potential ground movement to be provided.

Table 1.1. Monitoring locations and devices.

Location	Inclinometers	Acoustic Inclinometer	Piezometers	Weather station
Runswick Bay	4			
Whitby West Cliff	1			
Robin Hood's Bay	2		4	
Scalby Ness	4		14	
Scarborough North Bay – Oasis Café	2		3	
Scarborough North Bay – The Holmes	2		6	
Scarborough South Bay	17*	1	38*	
Filey Town	4		24	
Filey, Flat Cliffs	4	1	4	1
TOTAL	40	2	93	1

^{*}a single inclinometer and a diver piezometer with barometric diver was added at St Nicholas Cliff in 2014 between collection of the 1st and 2nd set of monitoring data.

1.3 Programme of work

The planned programme of future analysis and reporting is shown in Table 1.2, which assumes the final interpretative report will be provided three months following receipt of the preceding 6 months' monitoring data.

Table 1.2. Programme of data collection and reporting

JBA Monitoring Period	CH2M (Halcrow) Analysis Report
Data set 1: June 2012 to November 2013	Report 1: March 2014 (CH2M 2014a)
Data set 2: December 2013 to May 2014 (data received 1 Aug 2014)	Report 2: November 2014 (CH2M 2014b)
Data set 3: June 2014 to November 2014	Report 3: March 2015 (CH2M 2015a)
Data set 4: December 2014 to May 2015	Report 4: August 2015 (CH2M 2015b)
Data set 5: June 2015 to November 2015	Report 5: February 2016 (CH2M 2016a)
Data set 6: December 2015 to May 2016	Report 6: August 2016 (CH2M 2016c)
Data set 7: June 2016 to November 2016	Report 7: January 2017 (CH2M 2017)
Data set 8: December 2016 to May 2017	Report 8: October 2017 (this report)
Data set 9: June 2017 to November 2017	Report 9: February 2018
Data set 10: December 2017 to May 2018	Report 10: August 2018

1.4 Scope of data analysis work

JBA have sole responsibility for collection and checking of all inclinometer and piezometer data at 6 month intervals. JBA provide CH2M with the inclinometer and ground water data presented as graphs, ready for interpretation. The following graphs are provided in Appendices to this report:

- Inclinometer incremental displacement total displacement at 0.5m intervals down the length
 of borehole since the baseline reading along two axes (A0 being downslope, A180 being at right
 angles to the slope). This plot is free from errors associated with past readings as only the most
 recent and original readings are compared. This plot highlights the depths where most
 significant movement has occurred.
- Inclinometer cumulative displacement sum of all incremental displacements down the length
 of the borehole showing total deformation since the baseline reading along the two axes. If a
 user error has occurred, it is carried through all cumulative plots, potentially giving misleading
 results. Errors can usually be identified by comparison to incremental displacement plots.
- Inclinometer absolute position this plots the absolute position of the inclinometer casing when viewed vertically. While it does not give information on the rate of movement, it highlights the direction of any deformation and can be used to assess error in the data.
- Groundwater data from piezometer divers or data loggers these data are plotted as a continuous line showing groundwater level fluctuation relative to Ordnance Datum (OD).
- Groundwater data from monitoring wells these data are plotted as single points, showing
 groundwater level relative to OD at a particular point in time. They provide an independent
 check of piezometer data or water level information from boreholes that do not have automatic
 data logging capability.

The scope of CH2M's data analysis work involves the following tasks:

- Checks of inclinometer and piezometer monitoring data provided by JBA to ensure the correct information is provided, and identification of any obvious errors in the data.
- Downloading and analysis of meteorological data from the weather station installed at Filey Flat Cliffs and Scarborough Spa. The weather station at Filey Flat Cliffs was non-functional for several months in 2015 and therefore supplemental data has been purchased from the MetOffice for Filey, around 3km to the north-northwest. Since 11 January 2016 meteorological data from Scarborough Spa has been used because the reliability of the Flat Cliffs weather station remains poor.
- Acquisition of experimental acoustic inclinometer data from Loughborough University.
- Analysis and interpretation of the data, including commentary on short and long-term patterns
 of change and observed relationships between rainfall, groundwater levels and ground
 movement.
- Comment on the implications of the observed data with regard to cliff instability hazard and risk management, allowing SBC to take any appropriate action.

The following sections provide a site-by-site discussion of the history of cliff instability and the monitoring regime, and interpretation of the new monitoring data. Comment is made on the relationships between rainfall, groundwater and ground movement, and the implications of the new data with regard to cliff instability hazard and risk management.

1.5 Cliff instability hazard assessment

Cliff instability hazard at each monitoring location is presented using a simple colour-coding system that summarises the significance of the result (Table 1.3). The assessment provides a simple record of activity that will be developed in subsequent reports to indicate changing levels of hazard.

Table 1.3. Instability hazard assessment guidance level

Hazard (low to high)	Definition
Green	Situation normal. No change in groundwater level from previous records, which are low or falling. Movement in inclinometers within margin of error (<5mm).
Orange	Site requires attention. Moderate or large increase in groundwater level from previous records or moderate movement in inclinometers. Failure of equipment, unreliable or no data requires attention.
Red	Immediate action required. Significant movement of inclinometer indicating high cliff instability hazard potential. Carry out site inspection, consider increasing the frequency of monitoring and managing public access to the area.

1.6 Checks of monitoring equipment integrity

Following completion of checking and interpretation of the first round of monitoring in early 2014, several inclinometer readings appeared to be erroneous, with some locations showing potential ground movement. A series of checks were undertaken during 2014 to determine whether or not the data were accurate, the source of any errors, and the implications to cliff instability risk management. In most cases, the errors were systematic and represent minor settlement of the borehole casing that gives rise to a sinuous pattern of deformation. However, where random errors were reported, it is likely that the borehole is partially blocked, leading to the probe coming away from the key ways. The 17 potentially blocked boreholes were therefore repaired by means of high pressure water jetting that was undertaken in early 2015.

In all cases where systematic or random errors have been identified, it has been recommended that the current reading is taken as a new baseline against which future recordings are made. In this way, potentially misleading historical results leading to cumulative errors will be removed. However, in order to determine whether change has occurred in the preceding 6 month period, data are also compared to the original baseline.

Weather Summary

2.1 Introduction

A meteorological station that records wind speed and direction, air temperature, humidity, air pressure, rainfall and rainfall intensity every 15 minutes has been in place at Flat Cliffs, central Filey Bay, since 29 September 2011. The device was inoperative from September 2014 to July 2015 and therefore supplemental MetOffice rainfall data were acquired from recording station Filey No 2 (54.20395, -0.30127), c. 3km north-northwest of Flat Cliffs. The Flat Cliffs weather station again failed in the period March to May 2016, however at this time a new weather station at Scarborough Spa had become operational and therefore data from that site has been used from 11 January 2016 onwards.

Data from all three weather stations are summarised in Table 2.1 and Figure 2.1. The records for the last six months show that winter 2016/17 and spring 2017 were relatively dry when compared to past records, with the exception of June 2017 which was close to the long-term mean (upper range).

Table 2.1. Monthly rainfall recorded at Flat Cliffs or Scarborough Spa met station

				Rainfa	ll (mm)			
Month	Long-term mean (upper range)	2011	2012	2013	2014	2015	2016	2017
January	80	No Data	31	41	113 (84.2)	No Data (13.4)	84* [part month]	14.5*
February	60	No Data	8	38	96 (71.2)	No Data (44.8)	20.7*	21.1*
March	60	No Data	27	32	29 (40.4)	No Data (22.2)	53.9* [part month]	22.7*
April	60	No Data	96	4	26 (33)	No Data (15.8)	43.4*	17.8*
May	60	No Data	34	37 [part month]	59 (50.8)	No Data (81.4)	15*	22.4*
June	80	No Data	104	No Data	34 (61)	No Data (41.2)	23*	67.5*
July	60	No Data	70	No Data	70 (93.2)	20	14.9*	6.1*[part month]
August	80	No Data	45	38 [part month]	No data (108.2)	17	69.7*	
September	80	0.14 (Part month)	69	15	No data (17)	46	13.8*	
October	80	35	53	52	No Data (58)	29	15.4*	
November	80	15	78	25	No Data (70)	77.3	50.9*	
December	80	72	132	6	No Data (27.2)	76.9	6.4*	

Note: Data in brackets are from Filey No 2 station. Data marked * are from Scarborough Spa

Daily rainfall totals recorded by the Scarborough Spa weather station are presented in Figure 2.2, which shows peaks on 17 May, 8 June and 28 June. The combined dataset has been used for comparison with all coastal slope monitoring data in order to identify relationships. The data are taken to be representative of the whole Scarborough Borough Council frontage, but it is accepted that micro-climate effects may lead to local variations.

The Filey No 2 MetOffice data were provided as weekly totals and therefore the calculated totals do not precisely correspond to calendar months. The data show that the wettest month on record was December 2012 with 132mm, and that the wettest month during 2017 was June, with 68mm.

Seasonal totals are shown in Figure 2.3, which shows that the wettest seasons tend to be winter and summer and that the spring is the driest. The wettest season on record was winter of 2013/14 (i.e. December 2013, January and February 2014) that received a total of 244mm rainfall. The summers of 2012 and 2014 were also very wet, receiving 219mm and 211mm respectively. The winter of 2015/16 received around double the rainfall of 2014/5, but the spring was marginally drier. The winter of 2016/17 and spring 2017 have been the driest on record.

Wind speed and air temperature records from Flat Cliffs are presented in Figures 2.4 and 2.5. The data shows January 2017 as the windiest month in the monitoring period. Relatively gentle conditions persisted during the late-winter and spring 2017. Recorded temperatures during 2017 are average, except for the relatively warmer conditions during June 2017.

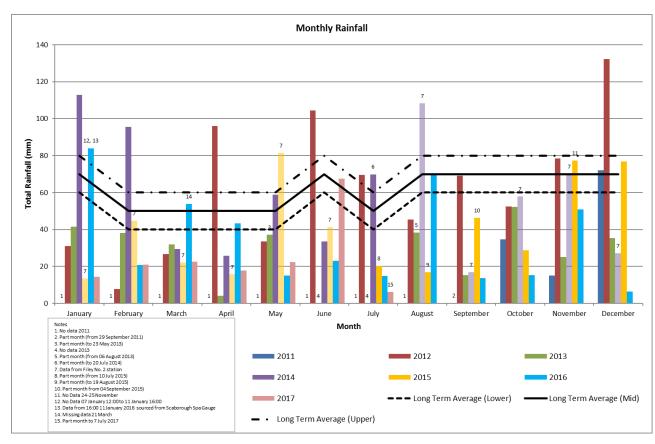


Figure 2.1 Comparison of monthly rainfall records (2011 to 2017).

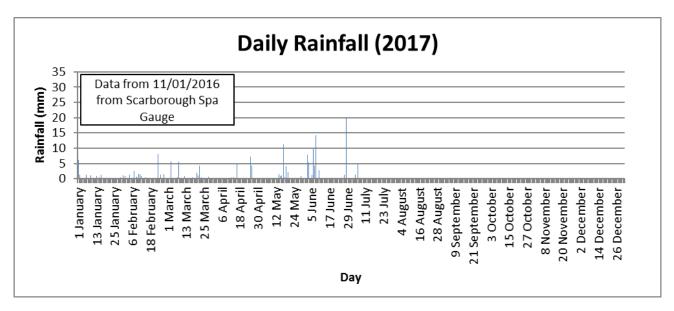


Figure 2.2 Daily rainfall recorded at Scarborough Spa during 2017

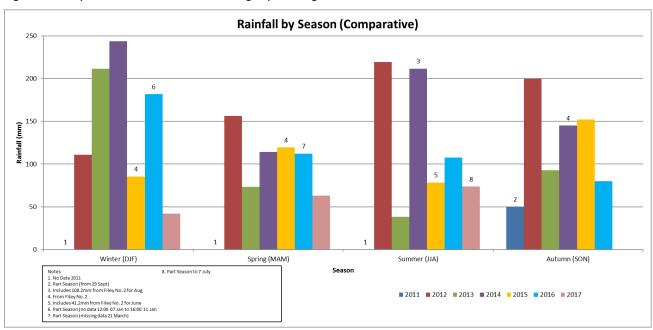


Figure 2.3 Seasonal rainfall comparison (2011-2017)

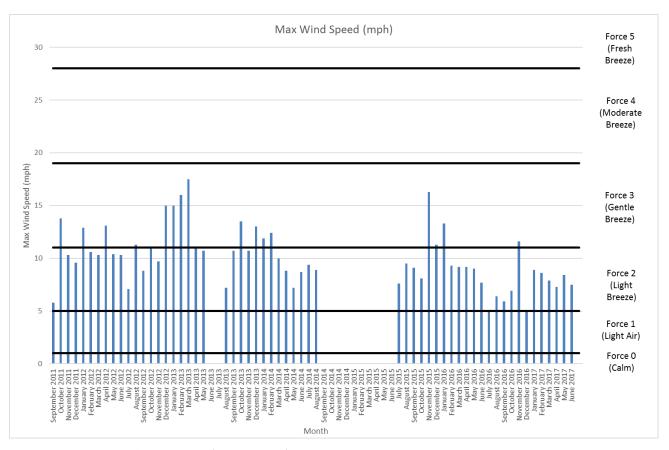


Figure 2.4 Maximum daily wind speed (2011 to 2017)

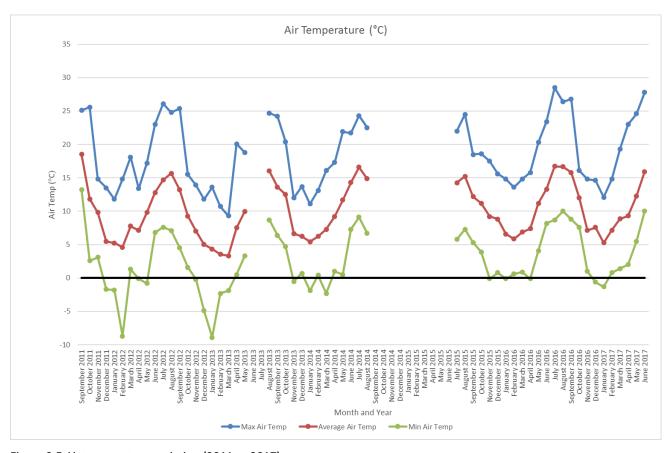


Figure 2.5 Air temperature variation (2011 to 2017)

2.1.1 Rainfall and landslides

The relationship between rainfall and the occurrence of landslides is complex and site-specific. It is often the case that a single intense rainfall event has little effect on a slope formed of relatively impermeable clay strata and soils, and instead cliff instability is only triggered after a period of sustained rainfall that allows groundwater levels to rise above a threshold level. This cumulative effect of sustained wet weather is known as antecedent rainfall. The time period over which antecedent rainfall exceeds a threshold for instability will vary from site to site, based principally on the local hydrogeology. It may vary from a period of days or weeks for sites formed of relatively higher permeability soils and rocks where groundwater responds rapidly to rainfall, to a period of months at locations of lower permeability soils and rocks.

The weather records for the SBC frontage spans a relatively short time period, but does include the particularly wet year of 2012. The only 'significant' ground movements at this time were recorded in BH7 at Scalby Ness, which occurred during December 2012. Monthly rainfall totals are provided in Table 2.1 and antecedent totals are presented in Figure 2.6. Assuming that rainfall was the sole trigger of this ground movement, it suggests a three-month antecedent rainfall threshold of 263mm is required to trigger movement. The absence of movements elsewhere on the coast at that time suggests that the antecedent rainfall threshold levels are above this at other locations.

Antecedent rainfall over the current December-May monitoring period show peaks that are significantly lower than that seen in December 2012 and with a similar pattern to the previous June to November monitoring period. It is therefore concluded there is a low likelihood of rainfall-induced landslides occurring in the monitoring period.

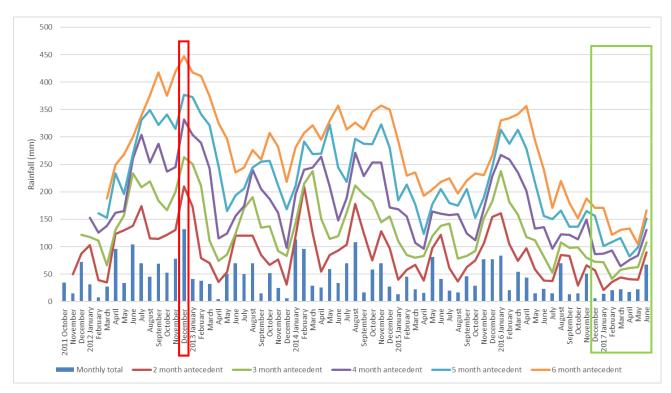


Figure 2.6. Monthly rainfall and two to six month antecedent totals (2011 to 2017). Ground movements were recorded at Scalby Mills during December 2012 (red box). Current monitoring period shown by green box.

2.2 Summary

The weather data collected to date highlights the following:

- 2012 was exceptionally wet, particularly in the months of April, June, July, November and December.
- 2013 was dry. After an unusually stormy spring period the temperatures remained high throughout the summer and rainfall in all months was below average.
- January and February 2014 were much wetter than average, and the period March to July 2014 was comparatively dry.
- While no data were recorded from early September 2014 to February 2015, a review of Met Office records shows the Autumn 2014 period was characterised by dryer than average conditions.
- MetOffice data purchased from Filey shows that the period Dec 2014 to April 2015 was generally much drier than average. Only May 2015 shows wetter than average conditions
- Data from Flat Cliffs collected in late 2015 shows September was wetter than average, and
 December was wet, although not exceptionally so. Rainfall peaks occurred on 14 September and
 21 November and a sustained period of wet weather occurred from 25 to 30 December.
- Scarborough Spa weather station data collected over 2016 has shown that January, March and April have been slightly wetter than average. Rainfall peaked on 3 January and 28 March.
 Overall, data has shown Dec 2015 to May 2016 to have been typically wet, with mild weather conditions.
- Between June and November 2016, rainfall has been lower than average apart from August where significant rainfall occurred on 4 and 25 August. Conditions over the 6-month period have been relatively dry and mild. Rainfall between December and May 2017 has been lower than the long-term average, however in June 2017 rainfall was above average, peaking on 8 and 28 June. Overall, data shows the 6-month period to have been relatively dry, with mild weather conditions suggesting a low likelihood of rainfall-induced landslides occurring.

Runswick Bay

3.1 Site description

Runswick Bay is the northern-most instrumented site on the Scarborough Borough Council coastline and is located 16 km north west of Whitby. The bay is formed in weak glacial sediments between the more resistant Jurassic-age bedrock headlands of Caldron Cliff to the north and Kettleness to the south. The village of Runswick Bay is developed on a coastal slope formed in glacial sediments and weathered shale bedrock and is bordered by incised valleys of the Runswick Beck and Nettledale Beck. The village and all existing monitoring devices are located in cliff behaviour unit MU7/1 (Figure 3.1).

The village has a long history of coastal instability, with records dating back to 1682 when the whole village was destroyed by landslides. It benefits from a coast protection and slope stabilisation scheme that was constructed in 2001-02 that comprises sections of seawall and rock armour together with drainage, piling and earthworks. The village has been subject to a strategy study review to improve the standard of protection of the coast protection measures and remedy minor issues with the 2001-02 scheme (Halcrow, 2016b). A Design and Build contractor for the scheme has been appointed and site works are due to begin in September 2017, with completion by May 2018.

3.2 Ground model and monitoring regime

The ground model for Runswick Bay was developed by High Point Rendel in the 1990s as part of the original strategy study for the area (High Point Rendel 1998). Their work included drilling a series of instrumented boreholes, geomorphological mapping and stability analysis. This work highlighted three landslide complexes that threaten properties and infrastructure:

- Topman End (MU7/1) steep till slopes (30° to 40°) between Nettledale Beck and continuing north to Runswick Beck. The village is sited on this landslide complex. The slopes are characterised by an extensive pattern of small scarps and tension cracks behind small shallow failures. Mid-way down the slope the profile shallows to between 5° and 10° over a distance of 10-15m. Where the slope angle exceeds 35° there are numerous shallow failures that tend to be caused by excessive water entrainment and generally leave behind triangular scars bounded by steep sides and disrupted vegetation. The mechanism is uncertain, but High Point Rendel (1998) suggests a model of superimposed mudslide lobes.
- Upgath Hill (MU 7/1) is the area north of Runswick Beck, beyond the village. The cliffs are formed in weathered Upper Lias shales capped by sandstone beds of the Saltwick Formation and thin veneer of till. Cliffs are fronted by steep talus slopes (20 to 30°) that are protected by a reinforced concrete sea wall. The toe of the southern facing slopes is continually undercut by stream flow in Runswick Beck. Over the years Runswick Beck has cut down through the weathered shale forming an incised valley with sides that are characteristically over-steep. The failure mechanism is believed to be rockfalls with shallow mudslides developed in the talus slope.
- Ings End (MU 7/2 and 7/3) comprises a series of sub-vertical head scarps, up to 2.5m in height, below the cliff top between incised valleys of Nettledale Beck and Limekiln Beck, south of the village. Movement here would adversely impact the village car parks and could trigger movement in Topman End. The headscarps front undulating, low angle slopes formed in till, characterised by springs, streams and water ponding. Shear surfaces are believed to be curved, suggesting the landslide is an ancient degraded multiple-rotational complex with superimposed shallow mudslides that are active during periods of prolonged heavy rainfall.

The monitoring regime at Runswick Bay comprises four inclinometers that are installed within piles of a portal frame shear-key system designed to stabilise the slope within the Topman End landslide (Figure 3.1). The inclinometers were originally intended to monitor the response of the piles to loading, but due to uncertainty over methods to achieve this, the data has been used to simply monitor ground movement and performance of the piles.

3.3 Historical ground behaviour

A summary of historical data, adapted from Mouchel (2012) is summarised in Table 3.1. Overall, the data show no ground movement since 2009 and only subtle variation in groundwater levels, and therefore no relationship between groundwater level and ground movement has been identified.

Table 3.1. Summary of historical ground behaviour at Runswick Bay.

Observations in Mouchel 2012 (covering 6-month period between Dec 2011 and June 2012)	Total change observed between July 2009 and June 2012
Slopes indicated as stable. Groundwater levels variable across site in inclinometers, with no change since previous reading, except for A002 that showed a marked drop in water level since Dec 2011.	5mm movement indicated in A001 between 22.0 and 20.0 metres depth and in A004 from 10.0m depth increasing to 15mm at 2.0m depth. Groundwater is relatively static in each borehole, although A002, A003 and A004 experienced lowering of levels in summer 2011, with recovery to previous levels by Dec 2011.

3.4 New data

All monitoring data at Runswick Bay is at the Topman End landslide, and is solely intended to monitor the effectiveness of the piles installed in the late 1990s to stabilise the slope. Water-levels within inclinometer tubes installed in the piles were recorded under the previous Mouchel contract. This has not been continued in the current phase of work as it was recognised that the data were of limited value to slope stability assessments and could be misleading. Inclinometer data are summarised in Table 3.2. These data indicate no movement in the piles.

3.5 Causal response relationships

No ground movements have been recorded at Runswick Bay over the monitoring period. Groundwater levels were previously monitored within the inclinometer tubes installed in piles, however, these data are unreliable, and no ground water monitoring is planned at this location. This means determining a relationship between rainfall, groundwater response and ground movement at Runswick Bay is not possible with the current monitoring set-up.

3.6 Implications and recommendations

There are no implications or recommendations arising for this site. Monitoring of the inclinometers should be continued to check the integrity and stability of the piles.

Table 3.2. Summary of inclinometer data at Runswick Bay

Summary of Borehole		Report status						status	_ Change Dec – May 2017
20.0.0.0	past data	1	2	3	4	5	6	7	change see may serv
A001	Data collected from within 22m deep concrete pile near the top of the slope. The data indicates no significant movement has been recorded in the pile							Incremental movements less than 1mm during the monitoring period, which is insignificant.	Incremental movements less than 1mm during the monitoring period, which is insignificant.
A002	Data collected from within 17m deep concrete pile near the top of the slope. The data indicates no significant movement in the pile.							Incremental movements less than 1mm during the monitoring period, which is insignificant.	Incremental movements less than 1mm during the monitoring period, which is insignificant.
A003	Data collected from within 10.5m deep concrete pile near the bottom of the slope. The data indicates no significant movement in the pile.							Incremental movements less than 1mm during the monitoring period, which is insignificant.	Incremental movements less than 1mm during the monitoring period, which is insignificant.
A004	Data collected from within 10.5m deep concrete pile near the bottom of the slope. The data indicates no significant movement in the pile up to Dec 2011.							Incremental movements less than 1mm during the monitoring period, which is insignificant.	Incremental movements less than 1mm during the monitoring period, which is insignificant.

Whitby West Cliff

4.1 Site description

Whitby West Cliff extends from the West Pier of Whitby harbour to Upgang Beach and Sandsend (Figure 4.1). A short (c. 500m long) section at the eastern-most extent fronting the Whitby Spa Complex comprises Jurassic-age limestone, sandstone and mudstone of the Scalby Group overlain by glacial sediments (CBUs 11/3 and 11/4), but the greater part of the cliff line is cut entirely in glacial sediments (CBUs 11/1 and 11/2). The cliffs cut in glacial sediments have a long history of instability and numerous relict landslide scars associated with shallow failures and seepage lines are visible. West Cliff benefits from coastal defences and slope stabilisation measures comprising a seawall, slope drainage and slope re-profiling that were installed in phases between the 1930s and 1970s. These measures have significantly reduced the risk of cliff instability, but they are near the end of their design life and distress in the slope has been observed.

4.2 Ground model and monitoring regime

The cliff instability features of West Cliff comprise shallow mudslides that are periodically active, but there is a concern that deep-seated failures may develop. The defended stretches show evidence of historical failures and despite toe protection the slopes are susceptible to periodic phases of movement associated with sustained rainfall. The unprotected cliff sections at Upgang beach have active mudslides. Historically, the monitoring regime at Whitby West Cliffs has comprised a series of survey pins that follow the line of the slope, which were intended to record deformation associated with cliff instability, and a single inclinometer (BH2) located near the base of the slope to the west of the Whitby Spa complex within CBU 11/2 (Figure 4.1). The inclinometer was read at 6 monthly intervals and also dipped to record water level. Survey pin data revealed no significant change during the period of monitoring by Mouchel. As water-level data derived from inclinometers is not recommended and liable to error, these readings are no longer taken and the current monitoring regime comprises six-monthly inclinometer readings only.

4.3 Historical ground behaviour

A summary of historical data, adapted from Mouchel (2012) is summarised in Table 4.1. Overall, the data show no deep ground movement since 2009 and only subtle creep of the upper metre of the slope, which is typical of glacial sediments. Groundwater data collected by dipping the inclinometer tube appeared to show a relationship with tide level and not groundwater. Groundwater data collected in this way are known to be very unreliable and therefore no relationship between groundwater level and ground movement can been identified.

The single monitoring location means the data from BH2 may not be representative of all of West Cliff. Caution should therefore be taken before extrapolating results across the site and monitoring should be supplemented with regular site inspection.

Table 4.1. Summary of historical ground behaviour at Whitby West Cliff

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)	Total change observed between July 2009 and June 2012
Survey pins show a total of 3mm movement at ground surface. Inclinometer indicates local slopes are stable, with surface creep in the top metre of ground.	Survey pins show -7mm movement in the top metre of ground. Inclinometer indicates local slopes are stable.

4.4 New data

Current data from the single inclinometer installed at Whitby West cliff is documented in Table 4.2 below.

Table 4.2. Summary of inclinometer data from Whitby West Cliff

Borehole	Summary of past data				F	Repo	rt st	atus	Change Dec to May 2017		
		1	2	3	4	5	6	7			
вно2	Inclinometer installed in a 20m deep borehole that passes through glacial sediment. Ground level is 13.78m OD and the base of the borehole is at -6.22m OD.							Incremental movements less than 1mm during the monitoring period, which is insignificant.	Incremental movements less than 1mm during the monitoring period, which is insignificant.		

4.5 Causal-response relationships

No relationships have been detected at this location.

4.6 Implications and recommendations

Monitoring at Whitby West Cliff is limited to a single inclinometer located near the base of the cliff to the west of the Whitby Spa complex. The device has not highlighted any cliff instability within the glacial sediments, although shallow failures have been observed on the cliff face during regular walkover inspections. The absence of any water level data at Whitby means it is not possible to determine the relationship between rainfall and ground movement, therefore, opportunities for installation of automated piezometer(s) should be considered.

Robin Hood's Bay

5.1 Site description

Robin Hood's Bay village is located on the coastal slopes and cliff top area of the northern-most part of Robin Hood's Bay. The cliff top part of the village is known as Mount Pleasant. The old village, situated on the coastal slope, has a long history of landsliding and currently benefits from a coast protection and slope stabilisation scheme that was installed in 2001.

The area being monitored in this study is the Mount Pleasant area, between Victoria Hotel and the cliffs to the north, where cliff instability is a concern. Cliff behaviour units in this area are composite cliffs formed of near-vertical sea-cliffs cut in Lower Jurassic clays overlain by glacial sediments. CBU 16/1 fronts Mount Pleasant and CBU 16/2 fronts the Victoria Hotel and the slope down to the old village (Figure 5.1). This section of coastline is not defended and has no slope stabilisation measures. Despite the bedrock cliff eroding at a slow rate, the overlying glacial sediments are prone to instability, and landslides occur episodically in response to sea cliff erosion and/or prolonged wet weather.

5.2 Monitoring regime

In response to the risk from landslides affecting the village, four instrumented boreholes have been installed in CBUs 16/1 and 16/2. These comprise two inclinometers and two double piezometers installed in bedrock and glacial sediments (Figure 5.1).

5.3 Historical ground behaviour

Robin Hood's Bay was not included in the original programme of monitoring and the first readings were taken in March 2010. The readings documented by Mouchel (2012) are summarised in Table 5.1.

Table 5.1. Summary of historical ground behaviour at Robin Hood's Bay

Observations in Mouchel 2012 (covering 6-month period between Dec 2011 and June 2012)	Total change observed between July 2009 and June 2012
Inclinometer BH2 shows movement at 22m depth. BH4 shows movement at 25m depth. Groundwater levels reduced.	n/a. First investigated in Dec 2011. Total change is as recorded between Dec 2011 and June 2012.

5.4 New data

The inclinometer and piezometer data recorded up to May 2017 is summarised in Tables 5.2 and 5.3.

Inclinometer data shows no significant movements recorded at borehole BH4. No data is available for BH2 due to the borehole being obstructed at the time of the site readings.

The piezometer data show groundwater levels have remained relatively steady or fallen over the monitoring period. BH1a, which is a shallower piezometer, shows groundwater levels have returned to previous levels observed following removal of tarmac that had been mistakenly laid over the boreholes. The deeper borehole BH1b is dry and should be checked as equipment may be damaged and requires attention to determine whether they can be repaired. Readings for piezometer BH3a shows groundwater levels have fallen significantly during the monitoring period, whilst the nearby deep borehole BH3b is dry and its integrity should be checked.

Table 5.2. Summary of inclinometer data from Robin Hood's Bay

Borehole	Summary of past data					F	Repo	ort status	Change Dec to May 2017
		1	2	3	4	5	6	7	
BH2	The borehole is 41m deep but inclinometer records are only provided for the upper 22m. Ground level is c. 55.1m OD. The recorded pattern of movement is hard to explain and is likely to represent accumulated error.							Incremental movements less than 1mm during the monitoring period, which is insignificant.	No data available, borehole obstructed.
BH4	The borehole is 40m deep and passes through 12m of glacial sediment and 28m of siltstone bedrock. Ground level is c. 74.2m OD and the base of the hole is at 34.2m OD. Cumulative movement plots suggest error in the data.							Incremental movements are less than 1mm since the last reading. From 30m depth 2mm minor displacement.	Incremental movements less than 1mm during the monitoring period, which is insignificant.

Table 5.3. Summary of groundwater data from Robin Hood's Bay

Borehole	Summary of past	Groundwater summary				Re	po	rt s	status	Change Dec to May 2017
Borenoie	data	Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017
BH1a	Ground level is 51.63m OD, the piezometer tip is targeting a shallower horizon. Water-levels have remained reasonably constant at c. 30m OD since installation.	22.7m OD 39.7m OD 17m							Decrease in groundwater level to 29.4m OD, returning to previous levels observed before the tarmac covering which likely affected previous readings.	Groundwater level at 29.4m OD, remaining stable around levels previous observed before the covering of tarmac, which may have affected previous readings.
BH1b	Ground level is 51.63m OD, the piezometer tip is targeting a deeper horizon. Water levels in this elevation have been less variable, having remained at around 37.6m OD.	37.6m OD 39.9m OD 2.3m							Readings increase by 1m to 38.8m OD, well within historical range.	Borehole dry. Check piezometer integrity.

Borehole	Summary of past	Groundwater summary				Re	po	rt s	tatus	Change Dec to May 2017
Dorenoie	data	Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017
внза	Ground level is 60.35m OD, the piezometer tip is targeting a shallower horizon. Water level has remained between 44.3m and 44.8m OD between installation in March 2010 and May 2012.	44.5m OD 56.1m OD 11.6m							Site readings lost due to data logger saving error. Readings should be re-taken on next visit.	Readings fall from 56.2m OD in October 2016, close to historical low at 47.5m OD in May 2016, the lowest groundwater level since 2012.
BH3b	Ground level is 60.35m OD, the piezometer tip is targeting a deeper horizon. Water levels have fluctuated by no more 2m about a mean of c. 56m OD. Low groundwater levels occurred in May 2010 and highs occurred in July 2010 and Nov 2011.	47.5m OD 56.7m OD 1.4m							Site readings lost due to data logger saving error. Readings should be re-taken on next visit.	Borehole dry. Readings from previous monitoring period indicate groundwater levels had fallen to historical low of 47.5m OD in October 2016. Check piezometer integrity.

5.5 Causal-response relationships

A subtle relationship between rainfall and groundwater levels, particularly in the shallower piezometer BH1a, was observed for the wet December of 2011 and the wet summer of 2012, and wet winter of 2015/2016. However, the dry conditions of 2013 were not reflected in the groundwater data, suggesting surcharge of groundwater from local sources may be occurring. Water levels in BH3a have fallen significantly in 2017 to their lowest since 2012, which may reflect the exceptionally dry conditions during winter 2016/17 and spring. There is also the possibility that the low resolution of monitoring at this location, particularly in shallow piezometers, may simply be picking-up short duration responses to brief but intense rainfall events.

5.6 Implications and recommendations

The groundwater data indicates a continuation of past patterns at Robin Hood's Bay. BH1a shows groundwater level has significantly fallen to previous levels observed before the tarmac covering, and remain steady. BH3a indicates groundwater has fallen significantly. These locations should be checked and the next monitoring data reviewed, whether this trend continues. Boreholes BH1b and BH3b were dry should be checked on the next site visit.

Previous work by Mouchel has noted that piezometer tubes have progressively become shallower, suggesting ingress of sediment. It is therefore recommended that the piezometer tubes be flushed out. Results from inclinometers are hard to interpret, meaning there is uncertainty over the nature of any recent ground movement. These data should be carefully reviewed in future monitoring reports and erroneous data removed from record. To improve understanding of the relationship

between groundwater and rainfall, this site would benefit from installation of automated piezometers to provide a continuous record of groundwater fluctuations.

Scalby Ness

6.1 Site description

Scalby Ness is the promontory that forms the northern boundary of Scarborough's North Bay. The headland is incised by Scalby Beck which flows through a steep-sided valley cut in glacial sediments and the underlying Jurassic sandstone/siltstone bedrock. Scalby Beck acts as a flood relief channel for the River Derwent via the 'Sea Cut', a man-made channel connecting the Derwent with the headwaters of Scalby Beck. The south side of the beck has housing that is threatened by ground instability in the over-steepened slopes cut in glacial sediments.

6.2 Ground model and monitoring regime

This site includes the cliff behaviour units MU19/11 and MU20/1 (Figure 6.1). The strategy study into the instability problems (Halcrow, 2005) characterised the area into three distinct landslide systems:

- CBU1 (northwest slopes) periodically active translational landslides in glacial sediment that lead to gradual headscarp recession. Instability is partly caused by toe erosion by Scalby Beck, but rising ground water levels following prolonged or intense rainfall are the principal trigger.
- CBU2 (northern part of the northeast slopes) large, ancient, deep-seated, periodically active
 landslide. Back-tilted blocks indicate a rotational failure, but translational mechanisms are also
 possible. Instability is partly caused by toe erosion by Scalby Beck but rising ground water levels
 following prolonged or intense rainfall are the principal trigger.
- CBU3 (southern part of the northeast slopes) stable slopes that have been reprofiled when the Sealife Centre access road was constructed.

Both CBUs 1 and 2 are at risk of failure, particularly if groundwater levels rise significantly. CBU3 is not considered to be at risk.

The monitoring regime at Scalby Ness is summarised in Figure 6.1. The slope is instrumented with three inclinometers and fourteen piezometers, seven of which are automated. Two inclinometers and nine piezometers are on the slope itself and the remaining installations are positioned on the cliff top.

6.3 Historical ground behaviour

Ground movement and groundwater levels were monitored by Mouchel from July 2009 to June 2012 and limited additional records of groundwater data back to June 2004. Mouchel's observations showed significant movement in BH7 between June and December 2010. No relationship between groundwater level and ground movement was reported by Mouchel, although relationships between rainfall and ground water levels in piezometers with shallow tips are identified. The readings documented by Mouchel (2012) are summarised in Table 6.1.

Table 6.1. Summary of historical ground behaviour at Scalby Ness.

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)

Total change observed between July 2009 and June 2012

Mouchel's piezometer graphs show notable increases in groundwater level in some piezometers (WS4 and WS6) to May 2012.

Ground movement reported at 12.0m BGL in BH7 at contact between gravelly sand and sandstone between June and December 2010, indicative of a developing shear plane although this movement has not yet manifested itself as recession of the headscarp. A failure was observed near the base of CBU1 between March and April 2010.

They report decreasing groundwater levels in CBU1, and peaks in groundwater levels in the shallower piezometers linked to intense rainfall events. Deeper piezometers remained at approximately the same level and were therefore less susceptible to variations in rainfall

6.4 New data

Tables 6.2 and 6.3 summarise the monitoring data from the inclinometers and piezometers at Scalby Ness.

Table 6.2 Summary of inclinometer data from Scalby Ness

Borehole	Summary of past data					Rep	ort s	tatus	Change Dec to May 2017
		1	2	3	4	5	6	7	
L1(C003)	Borehole is c.32m deep and situated on the cliff top above CBU1. Ground level is 35.47m OD and the borehole extends to c. 2.5m OD. It passes through 29m of glacial sediment and 3m of sandstone/mudstone bedrock. No displacements of the inclinometer tube greater than 2mm.							Incremental movements are less than 1 mm during the monitoring period, which is insignificant.	Incremental movements are less than 1 mm during the monitoring period, which is insignificant.
L2(C002)	Borehole is c. 35m deep and situated on the cliff top above CBU2. Surface elevation is 34.1m OD and borehole extends to c1.0m OD penetrating c. 31m of glacial sediment and 4m of mudstone bedrock. No displacements of the inclinometer.							Incremental movements are less than 1 mm during the monitoring period, which is insignificant.	Incremental movements are less than 1 mm during the monitoring period, which is insignificant.
L3(C004)	Borehole is ca. 17m deep, surface is 13.4m OD and borehole extends to							No significant movement since last reading, except for minor	No significant movement since last reading, Minor displacement that extends to ca. 2m BGL in clay is

8.5m of g sediment mudston sandston weathere upper 3m Cumulati	e and 8.5m of e and e that is ed in the n. ve plot is ertical with otion of a harman		displacement, likely relatively shallow surface creep that extends to ca. 2m BGL in clay.	likely to be relatively shallow surface creep
and Dece	lacement) surface, due to			
deep and the mid-s CBU2. Su elevation OD and the extends to OD through glacial se 7.5m of s /mudston The cumus shows and displacen between and June the contas and ston and grave c.4.7m O Subseque show posnegative	rface is c. 16.7m he borehole to c3.8m igh 13m of diment and tandstone ne bedrock. ulative plot ound 20mm nent Feb 2011 2011, above act between e bedrock elly sand at D. ent readings sitive and		Incremental movements are insignificant at less than 1mm. There has been no additional movement along the shear surface at c. 11 to 12m depth. The minor displacement extending to c. 2m depth is relatively shallow surface creep.	Incremental movements are less than 1 mm during the monitoring period, which is insignificant. No further movement along shear surface at c. 11 to 12m depth. Very minor displacement up to c. 2m depth within sand associated with shallow surface creep.

^{*}Surface elevations and borehole depths calculated from digital elevation model

Table 6.3. Summary of groundwater data at Scalby Ness.

Borehole	Long-term Pattern	Groundwater					Re	Change Dec to May		
		summary Min/Max/Range	1	2	3	4	5	6	7	2017
P1a	Automated piezometer. Tip at appox.25.65m OD*. Surface elevation at c. 35.6m OD* (cliff top above CBU 1, co-located with P1b). Fluctuates between 27.5 and	27.1m OD 28.9m OD 1.8m							Data not downloaded since October 2015. Collect on next site visit.	Data not downloaded since October 2015. Collect on next site visit.

P1b	28.5m OD, with rapidly rising and falling peaks linked to higher rainfall and subsequent dry periods. Automated piezometer. Tip at c. 18.1m OD*. Surface elevation at c. 35.6m OD (colocated with P1a). Relatively steady ground water level at ca.18.5m OD although fluctuations up to ca. 19.0m OD occur.	18.4m OD 19.2m OD 1.8m				Data not downloaded since October 2015, to collect on next site visit.	Data not downloaded since October 2015. Collect on next site visit.
P2a	Automated piezometer. Tip at c. 25.6m OD*. Surface elevation at c. 34.7m OD* (colocated with P2b). Fluctuates between 27.5 and 28.5m OD with peaks overlying a general trend of increasing water. Peaks and general trend correspond to the Filey rainfall record.	27.3m OD 28.7m OD 1.4m				Groundwater levels falling gradually to 27.5m OD.	Groundwater levels rise to 28.3m OD in March 2017, and fall gradually to 27.7m OD by May 2017.
P2b	Automated piezometer. Tip at c0.6m OD*. Surface elevation at c. 34.7m OD* (colocated with P2a). Prior to Oct 2009, ground water levels appear generally steady at c. 1.2m OD, except for fluctuations up to 2.5m OD in late 2007/early 2008. Records are absent between Oct 2009 and Mar 2010, after which levels are steady at around 2.5m OD.	0.9m OD 3.5m OD 2.6m				Readings remain steady at 2.4m OD.	Readings remain steady at 2.4m OD.
P3	Automated piezometer. Tip at c. 10.5m OD*. Surface elevation at c. 30.7m OD. Steady at around 14.6-14.7m OD until Oct 2009.	14.2m OD 17.5m OD 3.3m				Data not downloaded since May 2016. Collect on next site visit.	Data not downloaded since May 2016. Collect on next site visit.

P4a	Apparent recalibration between Oct 2009 and Mar 2010 after which groundwater levels are again steady at ca.17.2- 17.3m OD Automated piezometer. Tip at c. 8.3m OD*. Surface elevation at 18.6m OD (co-located with P4b). Fluctuating pattern occurs between June 2004 and Feb 2009 varying around 12m	12.7m OD 15.1m OD 1.4m				Groundwater levels steadily fall to 13.4m OD in saw-tooth pattern following previous readings.	Groundwater levels rise to 13.9m OD in March and gradually fall to 13.3m OD by May 2017 in a saw-tooth pattern following previous readings.
	to 13.6m OD. Peaks show steep rise and gentler fall, which is a characteristic response to heavy rainfall.						
P4b	Automated Piezometer. Tip at c. 6.35m OD*. Surface elevation at c. 18.6m OD (co- located with P4a). Fluctuating pattern between June 2004 and Feb 2009 with lows at around 12m OD and peaks to 13.6m OD. Peaks show steep rise and gentler fall characteristic of response to heavy rainfall	12.4m OD 14.8m OD 1.4m				Same pattern as P4a, but offset by c0.3m since early 2010. Piezometer may need recalibration.	Same pattern as P4a, but offset by c0.3m since early 2010. Piezometer may need recalibration.
WS4	Tip at 9.9m OD. Surface elevation at 16.3m OD (midslope, CBU 2). Fluctuations from c. 10m OD to c.15m OD in response to long-term/seasonal rainfall patterns. Limited response to short-lived rainfall peaks.	10.0m OD 15.4m OD 5.4m				Decrease in groundwater level by 1.6m to 13.8m OD.	Groundwater level falls by 1.3m to 12.5m OD.
WS5	Tip at 6.5m OD. Surface elevation at 11.3m OD (lower slope, CBU 2). Fluctuates between 6.5m OD and 7.5m OD between	6.5m OD 9.7m OD 3.2m				Borehole no longer functioning.	Borehole no longer functioning.

	September 2010 and June 2011 (low in summer/early autumn, high in winter).						
WS6	Tip at 9.72m OD. Surface elevation at 16.2m OD (midslope, CBU2). After an initial sharp rise post installation from ca. 10m OD to 12.5m OD, measurements from this piezometer show a gradual and uninterrupted increase to a high of 14.3m OD in May 2012.	10.0m OD 14.3m OD 4.3m				Groundwater levels remain steady at 13.3m OD.	Groundwater levels remain steady at 13.4m OD.
В6	Tip at 10.0m OD. Surface elevation at 18.55m OD (midslope, northern edge of CBU2). Pattern of substantial fluctuation, usually between 14m OD and 17m OD, with the exception of major low in August 2008 when installation may have been almost dry (groundwater level ca. 10m OD).	10m OD 13.8m OD 3.8m				Groundwater levels fall to 10.4m OD.	Borehole dry. Check piezometer integrity.
В9	Tip at 9.25m OD. Surface elevation at 17.8m OD (upper slope, CBU2). Fluctuation between ca. 10.0m OD and 12m OD except for substantial peaks in January 2008 (13.8m OD) and May 2008 (13.4m OD).	9.8m OD 16.7m OD 6.9m				Groundwater levels fall by 1m OD to 14.5m OD.	Groundwater level rises by 0.5m to 15m OD.
Sn2a	Tip depth at c. 13.9m OD*. Surface elevation at 16.35m OD* (co-located with SN2b). Likely that past results for 2a and 2b confused or tip depth for Sn2a incorrect;groundwat	12.5m OD 13.3m OD 0.8m				Ground water levels fall to 12.8m OD.	Borehole dry. Check piezometer integrity.

	er elevations not possible for tip depth stated.						
Sn2b	Tip depth at c. 8.35m OD*. Surface elevation at 16.35m OD* (co-located with SN2a). Likely that past results for 2a and 2b confused or tip depth for Sn2a incorrect. Sn2b shows groundwater levels between 1.6m OD and 11m BGL during 2011 and 2012.	10.3m OD 12.8m OD 2.5m				Groundwater levels continue to increase to new historical high of 12.8m OD.	Groundwater levels fall by 2m to 10.8m OD.

^{*}Indicates approx. tip and surface elevations calculated from elevation from digital elevation model and known tip depth, rather than topographic survey

The new data indicate:

- No significant ground movements recorded in any of the inclinometers.
- Groundwater levels have fallen in almost all boreholes with the exception of borehole B9. No significant movements over the monitoring period were indicated at inclinometers nearby B9.
- Considerable fall in groundwater levels in borehole Snb2 by 2m, following increase in previous monitoring periods to historical high.
- Water levels recorded in boreholes P4a and P4b follow the same pattern but at slightly differing levels, and it is recommended the piezometer calibration be checked.
- Piezometers in boreholes P1a, P1b and P3 require checking and downloading of data on next visit
- Piezometers in boreholes B6 and Sn2a were dry and their integrity requires checking on next site visit.

6.5 Causal-response relationships

The majority of shallow piezometers at Scalby Ness closely reflect the pattern of rainfall. Following a dry start to 2012, the spring and summer were exceptionally wet and the latter half of 2012 was wet. Peaks in groundwater were recorded during April, May July and December 2012. Overall, 2013, 2014, 2015 were drier than average. Piezometers installed with data loggers recorded falling levels until December 2013, after which groundwater levels rise and peak in mid-late February 2014, before falling and stabilising at lower levels by late 2014. Groundwater levels are typically lower than average during 2017, except for May where levels peak following higher than average rainfall. The above average rainfall in December 2015 and January 2016 is reflected in rising groundwater levels at many of the piezometers, however these levels have been maintained during drier months that follow and have either fallen or remained steady from May 2016 to 2017.

Deeper piezometers have a longer lag between rainfall and groundwater response. Those with data loggers show a much more muted response.

The inclinometers in BH7 and L2 show significant periodic sub-surface movement. BH7 is the most pronounced and indicates movement on an existing shear surface in glacial sediments above sandstone bedrock. Movement occurred between November 2013 and March 2014, associated with a period of high groundwater levels (nearby piezometers P4a and P4b show elevated groundwater

peaking in mid-February 2014 at 13.5 and 13.8m respectively). Neither inclinometer recorded movement between June and November 2014, associated with low groundwater levels. The precise relationship between groundwater level and ground movement is unclear. While movement in the winters of 2010/11 and 2013/14 can be associated with elevated groundwater, similarly high groundwater levels in the winter of 2012/13 and 2015/2016 are not associated with ground movement, possibly due to slow borehole equilibration with the surrounding ground.

6.6 Implications and recommendations

The groundwater data indicates levels have remained steady or fallen overall in the area. Rising water levels in Sn2b experienced from early to mid-2016 have returned to levels well below the historical high. The earlier rise in groundwater suggested a discharge from cliff top developments or a natural response to the localised movement at the slope toe, though no movements were recorded in adjacent inclinometers. In addition, piezometers P1a, P1b and P3 require readings to be re-taken on the next site visit. Boreholes B6 and Sn2a were dry and also require checking.

Scarborough North Bay - Oasis Café

7.1 Site description

Oasis Café cliffs are situated in the southern part of Scarborough's North Bay and occupy part of Clarence Gardens, which are landscaped coastal slopes open to the public (Figure 7.1). The cliffs rise to c. 30m OD and have a typical angle of 25-30°, although the main headscarp reaches 50°. The upper c. 15m of cliff is cut in glacial sediments and Jurassic sandstones and mudstones form the basal part of the cliff. The Holbeck to Scalby Mills strategy study (High-Point Rendel, 1999) classified the cliffs as multiple rotational landslides formed predominantly in the Jurassic bedrock. The landslides are fronted by the Marine Parade road and coast protection scheme and have not experienced toe erosion for over 100 years. Despite the toe protection, cliff instability risk in response to extreme rainfall remains a concern.

7.2 Ground model and monitoring regime

This frontage is covered by a single cliff behaviour unit, MU20/4a. Geomorphological mapping undertaken as part of the strategy study recognises a series of discrete landslides within this CBU, but all are classified as multiple rotational landslides formed predominantly in bedrock. It is assumed the basal shear surface is near Ordnance Datum and has formed in weak layers within the interbedded sandstones and mudstones. The monitoring regime comprises inclinometers and colocated automated piezometers at the cliff top, mid-slope and cliff toe positions aligned along a southwest to northeast bearing (Figure 7.1).

7.3 Historical ground behaviour

Table 7.1 summarises the observations in Mouchel (2012) from the monitoring undertaken at the Oasis Café.

Table 7.1. Summary of historical ground behaviour at Oasis Café

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)	Total change observed between July 2009 and June 2012
Static groundwater at around 8.05m at BH2p, and increase in water levels at BH3p and a decrease at BH4p. Slopes here appear to be stable from inclinometer readings although shallow ground movements were observed.	Apparent movements reported but these are attributed to operator error or temperature fluctuation rather than actual ground movements.

7.4 New data

Tables 7.2 and 7.3 summarise the monitoring data from inclinometer and piezometer installations at the Oasis Café

The new data indicate:

- No significant ground movements recorded in any of the inclinometers.
- Groundwater data show a continuation of, or further reduction in, recent low levels. This is in response to a drier than average winter and spring period. Short lived peaks in rainfall occurred in January and March and were recorded in BH3p, with a short lag.
- Elevated and fluctuating water levels in BH4p over this dry period are atypical and may reflect local sources of groundwater from cliff top properties.

Table 7.2. Summary of inclinometer data at Oasis Café

Borehole	Summary of past data					Rep	ort s	tatus	Change Dec to May 2017
		1	2	3	4	5	6	7	
вн3	BH3 is situated in the midslope and extends to c. 5.5m BGL. Surface elevation is 17.8m OD and the base of the hole is at c. 12.3m OD. The borehole extends through c. 3 m of glacial sediment before encountering 2.5m of mudstone, the uppermost metre of which is weathered. Past readings show no significant ground movement.							Readings are less than 1mm and therefore not significant.	Readings are less than 1mm and therefore not significant.
вн4	BH4 is situated on the cliff top and extends to ca.13.5m BGL. Ground level is 31.1m OD and the borehole extends to c 17.6m OD, penetrating 14m of glacial sediment and 3.5m of sandstone bedrock. Past readings show no significant ground movement.							Readings are less than 1mm and therefore not significant.	Readings are less than 1mm and therefore not significant.

7.5 Causal-response relationships

The winter 2013 to summer 2014 monitoring period was characterised by higher rainfall compared to the previous 6 months. The latter half of 2014 and 2015 were slightly drier than average and water levels tend to show very slight falls with superimposed monthly fluctuations. The higher than average rainfall in early winter 2015/2016 is reflected by elevated groundwater levels which fall in response to drier than average conditions which follow into 2017. The patterns seen in the past are still visible, with BH2p having an unclear response to rainfall and/or tides. Shallow piezometer BH3p shows a very rapid response to rainfall events (which probably explains the spikes on 10 Aug and 8 Oct 2014, and 9 May and 12 Dec 2015, 3 Jan and 27 August 2016), while only marginally deeper piezometer BH4p shows a lag response to prolonged periods of high rainfall. Groundwater levels in all boreholes remain below their peaks of winter 2012/13 and the inclinometers do not indicate movement.

7.6 Implications and recommendations

All the piezometers appear to read correctly and provide reliable data. The inclinometers also appear to be functioning correctly. No movements have been recorded at Oasis Café, and there are no specific recommendations at this location beyond on-going collection and analysis of data.

Future reports should pay particular attention to the midslope piezometer (BH3p) which shows rapid response to rainfall conditions, but no associated ground movements to date. In addition, attention

should be given to piezometer BH4p where rising groundwater over drier than average months may suggest an ingress of water from cliff top developments.

Table 7.3. Summary of groundwater data at Oasis Café

	Summary of past	Groundwater					Re	por	t status	Change Dec to May
Borehole	data	summary Min/Max/Range	1	2	3	4	5	6	7	2017
ВН2р	Tip depth at 8.05m OD. Situated in the lower cliff. Manual dip readings from Sept 2009 to May 2012 show fluctuation between 8.0 and 8.5mOD from Sept to Dec 2009 followed by no change to December 2011. Groundwater level then rises to 8.5m OD by May 2012.	7.9m OD 8.6m OD 0.7m							Continuation of past pattern. In October groundwater levels rapidly rise to historical high of 8.6m OD and fall to an average of c. 8.4m OD.	Continuation of past pattern, fluctuating weekly up to 0.5m. Groundwater levels falling in March 2017 to 8.1m OD and rise stabilising to average of 8.4m OD.
ВНЗр	Tip depth at 12.4m OD. Situated in the midslope. Manual dip readings from Sept 2009 to Dec 2011 show fluctuation between ca. 13.8m OD (June 2010) and 14.7m OD (Dec 2010). Final manual reading May 2012 shows substantial rise to 17.6m OD, reflecting high rainfall during spring 2012.	13.5m OD 16.7m OD 3.2m							Groundwater levels at an average of 14.5m OD. On 27 th August levels rapidly increased to 15.7m OD, coinciding with heavy rainfall on the 25 th . Groundwater levels rapidly returned to steady levels c. 14.2m OD.	Groundwater levels remain around 14.1m OD, peaking up to 15.9m OD in January and March, steadily falling to May 2017.
ВН4р	Tip Depth at 17.0m OD. Situated at the cliff top. Manual dip readings from September 2009 to May 2012 show groundwater levels fluctuating between 18.0m to 19.3m OD with peaks in April 2010, December 2010 and May 2012.	17.2m OD 19.4m OD 2.2m							Continuation of past cyclical pattern with sub-weekly variation, averaging 18.7m OD. Range of change small between 18.7 and 19m OD. Peaks in early October 2016. Groundwater level has decreased somewhat, but still remains at an elevated position.	Continuation of past cyclical pattern with sub-weekly variation, averaging 18.8m OD. Groundwater falls to 18.4m in late February, however through until May 2017 remains at a elevated position.

Scarborough North Bay – The Holms

8.1 Site description

The Holms is situated towards the southern end of North Bay, adjacent to Castle Headland. It is an area of sloping, hummocky, open parkland with a deeply-indented, arcuate headscarp between the castle at the cliff top and Marine Drive along the coast.

The slopes rise from Marine Drive at angles of c. 25-30° to a midslope bench at 35m OD and upper cliff at c.55m OD, where a near-vertical cliff face rises to the cliff top at c 85m OD. A variable thickness glacial sediments overlie interbedded sandstones and mudstones of Jurassic age. Two faults cross the site, one of which delineates the boundary of younger more resistant geological strata that form Castle Headland from the succession underlying much of the rest of North Bay.

The Holbeck to Scalby Mills strategy study (High-Point Rendel, 1999) classified the cliffs as multiple rotational landslides formed predominantly in the Jurassic bedrock. The landslides are fronted by the Marine Parade road and coast protection scheme and have not experienced toe erosion for over 100 years. Previous instability problems include a 200mm displacement of the sea wall, likely a result of reactivation of the pre-existing landslides. Movements of the main landslide body are estimated to be in the order of 10s of centimetres. Therefore, despite the toe protection, cliff instability risk in response to extreme rainfall remains a concern.

8.2 Ground model and monitoring regime

This site includes the Cell 1 cliff units MU21/1, which is the main landslide embayment, and MU20/4b which covers the cliffs to the west towards Oasis Café.

Mouchel (2012) state 'The Holms landslide system comprises 10 to 17m of landslide debris which overlies the intact Scalby Formation'. Two units within the landslide have been identified from ground investigations undertaken in 2000:

- An eastern unit, comprising a deep-seated landside which daylights close to the foreshore
- A western unit, composed of a shallower landslide which daylights approximately 1.5m above Marine Drive (c. 8.5m OD)

The monitoring regime at The Holms comprises:

- Lower slope two co-located piezometers. Each piezometer measures groundwater level at a different depth.
- Midslope two sets of two co-located piezometers, one set on the more north-easterly
 midslope bench and one set on the more westerly slopes. Each multiple piezometer location
 measures groundwater levels at different depths.
- Upper slope inclinometer in the central part, c. 50m NE and downslope of the bridge on the entrance road to the castle.
- Cliff top one inclinometer on the cliff top at the northern end of Mulgrave Place c. 50m to the west of the western end of the arcuate headscarp of The Holms.

8.3 Historical ground behaviour

The Holms was monitored by Mouchel between summer 2009 and summer 2012. A summary of their results is provided at Table 8.1. The pattern of groundwater variation at L1 appears to be affected by tidal influences and all other piezometers are affected by accuracy issues which prevent meaningful conclusions being reached about the groundwater regime at The Holms.

Table 8.1. Summary of historical ground behaviour at The Holms.

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)

Total change observed between July 2009 and June 2012

Mouchel (2012) comments that no ground movement has been indicated at BH10A. They mention continued ground movements of around 14mm between 13 and 10m depth (ca. 46-43m OD) in BH11. They report erratic groundwater readings from BH8 and BH9 a & b, and recommended flushing them as they believed they were blocked. As such, they report it was not possible to provide definitive information about the groundwater regime at The Holms.

Displacements of around 18mm at 10-13m depth (46-43m OD in BH11, 4mm of which occurred between December 2010 and June 2011 and a further 14mm between June 2011 and June 2012. Groundwater at L1 shows fluctuations of between 40mm and 120mm which is attributed by Mouchel (2012) to tidal level fluctuations.

8.4 New data

Tables 8.2 and 8.3 summarise the readings from the inclinometers and piezometers at The Holms up to May 2016.

The new data indicate:

- No significant ground movements recorded in any of the inclinometers.
- Groundwater data shows a continuation of past patterns. Water levels remained steady or decreased slightly, in response to the drier-than-average conditions over the monitoring period.
- Data collection now possible from piezometer BH9a. Piezometer BH9b is no longer functioning and should be investigated.

Table 8.2. Summary of inclinometer data at The Holms

Daniel al	Commence of most desired				F	Repo	rt St	atus	Change Dos to May 2017
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017
вн10А	BH10A c. 42m deep. Surface of borehole is 46.75m OD, base at 4.75m OD. Borehole passes 2m of made ground, 1m of clay and c.8m of clayey sand before encountering sandstone bedrock. Progressive movements in the positive A axis direction (upslope) are recorded between the surface and 5m BGL (a. 42m OD). The total maximum displacement that occurred by May 2012 was around 10mm.							Readings are less than 1mm and therefore not significant.	Readings are less than 1mm and therefore not significant.
BH11	BH11 is c.22m deep. Surface elevation is 55.86m OD, base at c.34m OD. Borehole passes through 5m of till before encountering weathered sandstone at c. 51m OD and intact sandstone at 41m OD. The inclinometer readings show a series of progressively larger deformations of around 20mm in the both axes within the weathered sandstone.							No change detected in sinusoidal pattern of deformation between 9 and 13m depth.	No change detected in sinusoida pattern of deformation between 9 and 13m depth.

8.5 Causal-response relationships

Rainfall has been lower than average since mid-2015, with the exception of above average rainfall in winter 2015/2016 and wet November 2016. The piezometers at The Holms show a lagged response to these conditions with only BH8a showing a rapid response to May 2015, March and November 2016 rainfall. Levels have since fallen during winter 2016/2017 following months of dry conditions. Other boreholes show a continuation of past fluctuating or steady levels of groundwater, suggesting they respond to several months' antecedent rainfall. Over the whole record, BH8b shows a different pattern of gradual highs followed by sharp falls, however movements are not shown in the inclinometer upslope at BH10A.

8.6 Implications and recommendations

There are no implications or recommendations arising for this site.

Table 8.3. Summary of groundwater data at The Holms

Borehole	Long-term Pattern	Groundwater					Re	poi	rt Status	
		summary Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017
L1a	Tip depth at -8.03m OD, co-located with L1b. Manual dip readings from June 2009 to May 2012 show limited groundwater with variation between 5.9m OD (June 2010) to 4.6m OD (March 10). Piezometer tip is deeper than BH1Lb, but shows a higher piezometric level that may indicate a confined aquifer under artesian pressure	0.5m OD 2.5m OD 2m							Continuation of past cyclical patterns, with 2 to 3 week variations of up to c. 0.3m in August 2016. The monitoring period represents an increase in average groundwater levels from 0.6m OD to 0.9m OD, which is very low compared to past records.	Continuation of past cyclical patterns, with 2 to 3 week variations of up to c. 0.5m in February. Groundwater levels show an average decrease throughout the monitoring period, from 1.3m OD to 0.6m OD. Groundwater levels peak in late December 2016 to c. 1.3m OD, and are lowest during late February at 0.5m OD.
L1b	Tip depth at -2.97m OD co-located with L1a. Manual dip readings between June 2009 and May 2012 show steady groundwater level around 1.9m OD.	3.7m OD 4.7m OD 1.0m							Continuation of 2 to 3-week cyclical pattern. Average groundwater level has fallen slightly at 4.2m OD. Groundwater level rapidly peaks and falls in early October 2016 to highs of 4.4m OD. Antecedent rainfall was low and therefore it is unclear what triggered this groundwater response, which may relate to ingress of water from cliff top developments.	Continuation of 2 to 3-week cyclical pattern. Average groundwater level has fallen from 4.3m OD in December 2016 to 4.1m OD in May 2017. Groundwater level falls close to the historical low at 3.8m OD in late February 2017, and rapidly rise again to steady levels. Antecedent rainfall was low and therefore it is unclear what triggered this groundwater response, which may relate to ingress of water from cliff top developments.

Borehole	Long-term Pattern	Groundwater					Re	epoi	t Status	
		summary Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017
вн8а	Tip depth at 10.16m OD. Borehole top at 31.16m OD Colocated with BH8b. Monitoring from Sept 2010 shows an initial fall in level to a low of 10.43m OD in June 2011 then a gradual rise to Dec 2011, reflecting wetter weather, before a sharp rise to 23.6m OD by May 2012 as a result of exceptional rainfall.	9.7m OD 10.7m OD 1.0m							Groundwater levels remain steady during the monitoring period, averaging 10.3m OD. A rapid rise and fall of 0.4m OD occurs in early October 2016, similar to other boreholes in the vicinity. Antecedent rainfall was low and therefore it is unclear what triggered this groundwater response, which may relate to ingress of water from cliff top developments.	Groundwater levels fall during this monitoring period from an average of 10.5m OD in Dec 2016 to 10.3m OD in May 2017. Large cyclical variation occurs between January and March by 0.5m, close to historical lows. Antecedent rainfall was low and therefore it is unclear what triggered this groundwater response, which may relate to ingress of water from cliff top developments.
BH8b	Tip depth at 3.16m OD. BH top at 31.16m OD, co- located with BH8a. Groundwater levels dropped from an initial high of 17.3m OD at installation in Sept 2010 to a low of 9.55m OD in Feb 2011. Levels then gradually rise through 2011 to c. 10.6m OD in Dec 2012 before a sharp rise to 22.2m OD by May 2012. This shows a similar rainfall-influenced pattern to BH8a.	9.4m OD 14.5m OD 5.1m							Groundwater levels continue to rise until August where they rapidly drop from 12.6m OD to 10m OD and continue to remain steady.	Groundwater levels continue a saw-tooth pattern, rising to 11.7m OD between Dec and Feb, then rapidly falling to 10m OD, remaining steady.
ВН9а	Tip depth at 9.49m OD. Surface at 33.49m OD colocated with BH9b. Shows sharp increase after installation from c. 11.5m OD to a high of 26.6m OD by Feb 2011 before falling to 24.3m OD in June 2011. Between June and Dec 2011 ground water levels rise again to around 27.0m OD before falling to 26.3m OD.	13.6m OD 26.2m OD 10.9m							No data recorded for monitoring period. Last data record for May 2016 at 19m OD. This site requires attention to fix or replace the piezometer and damaged cable.	Data retrieved for previous and current monitoring period shows groundwater falls until December 2016, after which levels rise in January to 18.7m OD. Levels remains steady at an average of 15.9m OD until May 2017.

Borehole	Long-term Pattern	Groundwater					Re	еро	rt Status		
		summary Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017	
вн9ь	Tip depth at 0.49m OD, surface at 33.49m OD colocated with BH9a. Shows sharp increase in ground water levels from c. 10m OD after installation in Sept 2010 to c. 25m OD in Feb 2011. Continues to gradually rise to c. 26m OD in June 2011 before gradual fall to 23.2m OD by May 2012. This pattern is similar to that recorded in BH9a, but contrary to that in BH8a and BH8b.	9.7m OD 30m OD 18.5m							Data retrieved for previous and current monitoring period shows gradual fall from peak historical high groundwater levels in January 2016 of 40m OD to October at 14.2m OD.	No longer functional. Piezometer should be repaired.	

Scarborough South Bay

9.1 Site description

South Bay is formed from cliffs cut in Jurassic sandstones and siltstones that are overlain by a thick sequence of glacial sediments. A series of deep-seated landslides have developed in the glacial sediments and underlying weathered bedrock in post-glacial times. Since Victorian times, the cliffs have been extensively landscaped into public areas that include the Spa conference centre complex. The coastline has marginal stability, but first time failures do occur: the Holbeck Hall Hotel landslide occurred in June 1993 and there are records of similar cliff failures occurring elsewhere along the frontage over the last several hundred years. The whole frontage benefits from coastal defences, but ground movements in pre-existing landslides and over-steep cliff sections continue to occur, particularly in response to periods of elevated ground water levels, and there remains concern of first-time failures and reactivation failures in the cliffs. Instability risk is therefore a concern along the whole of South Bay.

The majority of South Cliff (from St Nicholas Cliff to Holbeck Gardens) was mapped in 2011 as part of the Scarborough Spa Coast Protection scheme. This mapping underpins the ground model for this site. Cliff behaviour units (CBUs) have been defined and their activity status classified under the Cell 1 Regional Monitoring Programme.

9.2 Ground model and monitoring regime

Pre-existing landslides have developed in the thick sequence of glacial sediments that form the upper coastal slope. Their geomorphology generally comprises arcuate landslide embayments with mid-slope benches that are fronted by elongate mudslide tracks and vertical *in situ* bedrock cliffs. The basal shear surface typically appears at the contact between the glacial sediment and underlying Jurassic bedrock, but it is likely that the significant local variation in the glacial sediments allows secondary shear surfaces to form along clay layers.

The monitoring regime at South Bay is summarised in Appendix A and Figure 9.1. It comprises an extensive suite of inclinometers and piezometers, most of which are automated, and an experimental acoustic inclinometer installed near the Spa Centre.

The areas being monitored comprise, from north to south:

- St Nicholas Cliff till cliff fronting the Grand Hotel and cliff lift with a co-located single
 inclinometer and diver piezometer with barometric diver that were installed in 2014 (MU22/0)
- Spa Chalet Gardens till cliff with groundwater monitoring at its toe and an inclinometer inland
 of the cliff top (MU22/1).
- Spa Centre and gardens rotational landslide (MU 22/2) and very steep till cliff (MU22/3) in the
 vicinity of the Spa buildings. Extensive monitoring of groundwater levels and ground movements
 at locations inland of the cliff top, on the slope and at the cliff toe.
- Clock Café rotational landslide (MU 22/3) that is monitored with transect of devices comprising two inclinometers on the slope and a piezometer inland of the headscarp.
- South Cliff Gardens till cliff with a mudslide embayment north of the Rose Garden (CBU 22/5),
 a small rotational landslide at the Rose Garden and a much larger rotational landslide at the
 Italian Garden, known at the South Bay Pool landslide(CBU 22/6). The area is monitored by three
 transects of devices that cover each of the landslides.
- Holbeck Gardens (CBU 22/7) till cliff monitored a three locations.

These areas include both pre-existing landslides and also intact cliffs and headscarps where instability is considered to be a risk. The Spa Centre is the focus of monitoring and is also the subject of an on-going coast defence scheme to improve the seawall and stabilise the slope.

At each location a suite of instruments are installed on the promenade, on the coastal slope and at the cliff toe allowing ground models to be developed and stability modelling to be undertaken.

9.3 Historical ground behaviour

South Bay was monitored by Mouchel Ltd for the period between summer 2009 and summer 2012. A summary of their results is provided in Table 9.1, which shows slight movement in a number of inclinometers and variable groundwater levels. No relationship between groundwater level and ground movement was reported by Mouchel.

Table 9.1. Summary of historical ground behaviour at Scarborough South Bay.

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)	Total change observed between July 2009 and June 2012
AA10 (Clock Cafe) and AA08 (south Cliff Gardens) showed slight movement at shallow depths. Movement at greater depth was indicated in BHs 12, 13, 14 (at the Spa) and 16A (South Cliff Gardens). No movements indicated by other inclinometers. Groundwater levels are generally variable across the sites, except in the south of the Spa, where levels were reduced.	In addition to observations between Dec 2011 and June 2012, slight movement was recorded at AA04 in the upper 7m of ground, at AA10 in the upper 3.5m and at AA11 in the upper 3m. All net movements have been less than 10mm.

9.4 New data

For clarity, new data for South Bay are presented for each of the monitoring areas separately.

9.4.1 St Nicholas Cliff (MU 22A)

The cliff here is around 30m high and heavily landscaped with terraces and footpaths and formed in fine-grained glacial sediments (Figure 9.1A). Average slope angle is 20 to 30° but is locally steeper with sections supported by retaining walls. The cliff is crossed by a cliff lift and the cliff top is occupied by the Grand Hotel. There is no history of instability in recent years and this CBU was not reported on by Mouchel.

Table 9.2 Summary of inclinometer data at St Nicholas Cliff

Daniel ala	Community of many data				Repo	us		Change Dec to May 2017		
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017	
FR01	FR01 is situated above Foreshore Road in front of the Grand Hotel at 11.43m OD. The borehole is c.20m deep with its base at c8.5m OD and passes through c.10.5m of made ground and 9.5m of fine grained glacial sediments. FR01 has been monitored since 16 June 2014.	n/a						Readings show less than 1mm movement and are not significant.	Readings show less than 1mm movement and are not significant.	

Table 9.3 Summary of groundwater data at St Nicholas Cliff

Borehole	Summary of past data	Groundwater summary Min/Max/Range	1	2	3	Rep	ort st	atus 6	7	Change June to Nov 2016
FR02	FR02 has been monitored since 21 May 2014. Tip is at 18.0m depth (c6.5m OD). Pattern shows variation consistent with short and medium term tidal cycles.	7.4m OD 8.4m OD 9m	n/ a						Continuation of past cyclical pattern, with groundwater levels averaging at 7.6m OD.	Continuation of past cyclical pattern, with groundwater levels averaging at 7.6m OD.

No ground movement is recorded at this site and water levels are stable.

9.4.2 Spa Chalet (MU 22/1)

This cliff is very steep and formed in glacial sediment that does not appear to have been affected by landsliding. The cliff has been previously stabilised with soil nails and netting. Monitoring comprises a single inclinometer on the promenade and a pair of closely located piezometers at the cliff toe. Inclinometer data are summarised in Table 9.4 and piezometer data in Table 9.5.

Table 9.4 Summary of inclinometer data at Spa Chalet

Daniel ala	Commence of most data					Repo	ort Sta	Change Dec to May 2017		
Borehole	Summary of past data	1	1 2 3		4	5	6	7	Change Dec to May 2017	
BH12	BH12 is 65m deep (ground level at 48.05m OD, base at - 16.95m OD) and extends through 60m of glacial sediment and 5m of sandstone/mudstone bedrock. Cumulative readings show a pattern of subtle movement that is interpreted as error.							Readings are less than 1mm and therefore not significant.	Readings are less than 1mm and therefore not significant.	

Table 9.5. Summary of groundwater data at Spa Chalet.

Borehole	past data summary 1						Re	por		
	past data	summary Min/Max/Rang e	1	2	3	4	5	6	7	Change Dec to May 2017
BH12	Tip at -8.4 OD. Cyclical pattern with c. two- week frequency between peaks. Maximum levels are between 1.25 and 1.5m above OD and minimum levels are between 0.3 and 0.5m above OD. Given the tip is below mean sea-level it is possible the cyclical pattern is related to tides.	0.0m OD 2.3m OD 2.3m							No data recorded for monitoring period. Last data recorded in May 2016. This site requires attention, to fix or replace the piezometer and damaged cable.	No data recorded for monitoring period. Last data recorded in May 2016. This site requires attention, to fix or replace the piezometer and damaged cable.
BH12a	Tip at 3.6m AOD. High degree of variability, with rapid fluctuation about a mean water level of c. 3.6m above OD. Peak water levels are c. 3.9m AOD and minimum levels are c. 3.3m AOD.	3.2m OD 3.9m OD 0.7m							Range of fluctuations within past limits and linked to tidal cycles. Fluctuations range from 3.4 to 3.9m OD, averaging c. 3.6m OD. Groundwater rapidly peaks and falls in early October to 3.9m OD.	Range of fluctuations within past limits and linked to tidal cycles. Fluctuation range from 3.3 to 3.9m OD, averaging 3.6m OD. Large cyclical variation occurs between January and March 2017, close to the historical low.

No ground movement has been recorded and fluctuations in groundwater levels are within the ranges previously observed. The piezometer in borehole BH12 requires attention to fix or replace faulty equipment.

9.4.3 Spa (MU 22/2 and 22/3)

The Spa is the focus of monitoring in South Bay, with eight inclinometers and 21 piezometers installed in the area (Figure 9.1B). The cliffs are generally steep and formed in glacial sediment. Shallower cliff sections are associated with a deep-seated landslide seen immediately north of the Spa Centre and localised shallow landslides. The monitoring results are described in Tables 9.6 and 9.7.

Table 9.6. Summary of inclinometer data at the Spa

Banahala	Commence of weath data					Re	port	Status	Change Dec to May
Borehole	Summary of past data	1	2	3			7	2017	
AA04 (G2)	40.5m deep borehole penetrating 34.5m of glacial sediments and 6m of sandstone/siltstone bedrock. Ground level is 47.62m OD, base of hole is 7.12m OD.							Readings are less than 1mm and therefore not significant.	No significant movement.
BH13	61m deep borehole inland of the headscarp that penetrates 52m of glacial sediment and 9m of sandstone bedrock. Ground level is 53.93m OD, base of hole at -7.07 OD. Deflection of up to 80mm in the upper 35m (i.e. above 19m OD) of the borehole associated with creep.							No significant movements recorded since the last readings.	No significant movements recorded since the last readings.
BH14	55m deep borehole penetrating c. 50m of glacial sediments and 5m of sandstone bedrock. Ground level at 55.73m OD, base of hole at 0.73m OD. Uniform cumulative displacement of c. 5mm in the upper 35m of the borehole, with peaks of up to 10mm displacement from 35 to 55m depth. Readings are not progressive in time, suggesting shrink-swell behaviour.							No significant movements recorded since the last readings.	No significant movement.
BH101	Borehole is located in the seawall, beyond the toe of the Spa landslide and is 26.5m deep, passing through 21m of glacial sediment and 5.5m of sandstone and mudstone bedrock. Ground level is 6.77m OD and the base is -19.7m OD. No significant movement has been detected in the past.							No significant movement.	No significant movement.
ВН103	10m deep borehole that only penetrates glacial sediments. Ground level is 6.65m OD, base of hole at -3.35m OD. Apparent displacements between installation in Oct 2012 and Dec 2012 are <1mm.							No significant movement.	No significant movement.
ВН107	18m deep borehole that passes through 13m of glacial sediments and 5m of sandstone/mudstone bedrock. Ground level is 20.39m OD, base of hole at 2.39m OD. No displacements between installation in Oct 2012 and Dec 2012. Historical readings							Movements of less than 1mm in the borehole are not significant.	No significant movement.

	unavailable at current time therefore current reading cannot be compared to baseline.					
BH109	15m deep borehole that passes through 9m of glacial sediment and 6m of sandstone/mudstone bedrock. Ground level is 31.6m OD, base of hole is 16.6m OD. Apparent displacements between installation in Oct 2012 and Dec 2012 are <1mm.				Movements of less than 1mm in the borehole are not significant.	No significant movement.
BH105	45m deep borehole passing through 44m of glacial sediments an 1m of sandstone bedrock. Ground level is 41.75m OD and base of hole is -3.25m OD. Apparent displacements between installation in Oct 2012 and Dec 2012 are <1mm.				No significant movements since last reading.	No significant movement.
ВН105а	Acoustic inclinometer installed to a depth of 40m since 14 Nov 2012 adjacent to BH105. Ground level is 42m OD, base of hole is 2m OD. Since installation in Feb 2013, the device has detected a relatively low level of activity in response to rainfall events. No significant ground deformations have been indicated by the acoustic monitoring.				AE measurements between February 2016 and December 2016 do not show significant slope movements. As seen previously, periods of elevated AE activity are thought to be a response to rainfall events generating seepage in the gravel beds.	AE measurements between December 2016 and September 2017 do not show significant slope movements. As seen previously, periods of elevated AE activity are thought to be a response to rainfall events generating seepage in the gravel beds.

Table 9.7. Summary of groundwater data at the Spa

	Long-term	Groundwater	Report status						Change Dec to May	
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
H2a	Located near the headscarp of the Spa landslide. Tip at 17.3m AOD. 3 to 5 day frequency fluctuation around mean of c. 17.25m OD with amplitude of c. 0.5m. No clear long term trend or temporal pattern. Max water level 17.6m OD on 4 June 2013, min of 16.9m OD on 15 March 2013.	16.7m OD 17.6m OD 0.9m							No change in past pattern. Levels range from 17 to 17.5m OD. Groundwater levels peak in June and October 2016, suggesting a complex relationship to rainfall.	Groundwater levels reach the historical high at 17.6m OD in December 2016. Levels slightly decrease towards May 2017, averaging 17.3m OD, with large rises and falls in March close to the historical low of 16.8m OD.

	Long-term	Groundwater	Report status						status	Change Dec to May
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
H2b	Located near the headscarp of the Spa landslide. Tip at 11.1m AOD. 3 to 7 day frequency fluctuation around mean of c. 12.7m OD with amplitude of c. 0.3m. No clear long term trend or temporal pattern. Maximum water level 12.9m OD on 3 June 2013 and 7 July 2013, minimum of 12.3m OD on 14 December 2012.	12.0m OD 13.0m OD 1.0m							No change in pattern, groundwater levels steadily fluctuate around an average of 12.6m OD. During June 2016, groundwater levels do not fluctuate. Levels range between 12.5m OD to 12.7m OD.	No change in pattern, groundwater levels steadily fluctuate around an average of 12.6m OD. In March 2017, groundwater levels fluctuate greatest between 0.5m.
H5	Located near the base of the cliff behind the Spa building. Tip at 15.5m OD. Marked drop in water level from 22m OD in late 2012 to 17.5m OD in late 2013. Slight but shortlived recoveries on 5 Nov 2012 and 15 Aug 2013 when waterlevels rose by almost 1m in a day.	17.0m OD 23.01m OD 6.01m							Water levels rise to 22.2m OD between 19 and 29 June, then fluctuate between 21.5 and 22.4m OD for remainder of period. While water levels have remained high during this period, they have not reached historical highs of 23.1m OD. Note the new telemetry logger records water levels are 15 minute intervals, whereas the former system records hourly. Despite this higher resolution, the data recorded by the telemetry system show a smoother pattern, which suggests variations in earlier data are 'noise'.	Water levels follow a saw-tooth pattern, with a sharp rise and gradual fall that fluctuats between 22.2m OD and 22.7m OD. This pattern and range reflected the past data. The rapid rises in groundwater coincide with heavy rainfall in April and May, however some peak rainfall events recorded in the summer do not show the same response in groundwater levels. In the period since May the magnitude of change is greater and water levels are elevated between 21.6m OD and 22.7m OD. A sharp peak in groundwater level in February is assumed to be error.
1 spa	Located near the base of the cliff. Tip at 6.3m OD. Water levels fluctuate between c. 7m OD and c. 12m	6.7m OD 11.9m OD 5.2m							Groundwater levels fall to 7.8m OD in October 2016.	Groundwater levels rise to 9m OD in May 2017, though well below historical high.

	Long-term	Groundwater	Report status						Change Dec to May	
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	OD. High levels over 11m AOD occurred in May 2008, Dec 2009 to Apr 2009 with historical low of c.7m OD between Aug 2008 and Aug 2009.	,								
2 spa	Located near the base of the cliff. Tip at 6.4m OD. Water levels fluctuated between c. 10m OD and c. 12m OD between Jan 2003 and Aug 2009. Thereafter, variation increases with low levels recorded down to c. 8m OD. Low levels recorded during the winters of 2010 and 2011.	7.2m OD 12.1m OD 4.9m							Groundwater levels fall slightly to 10m OD in October 2016.	Groundwater levels rise slightly to 10.1m OD in May 2017.
3 spa	Located near the base of the cliff. Tip at 7.2m OD. As in '2 spa' water levels fluctuated between c. 12m OD and c. 13m OD between Jan until Aug 2009 and thereafter, variation increases with low levels recorded down to c. 7m OD.	7.1m OD 13.0m OD 5.9m							Groundwater levels fall slightly to 11.9m OD in October 2016.	Groundwater levels fall slightly to 11.8m OD in May 2017.
4 spa	Located near the base of the cliff. Tip at 10.9m OD. Very similar pattern to '3 spa'. Water levels fluctuated between c. 10m OD and c. 13m OD between Jan until Aug 2009 and thereafter,	6.1m OD 12.6m OD 6.5m							Groundwater levels fall to 11.6m OD in October 2016.	Groundwater levels fall slightly to 11.4m OD in May 2017.

	Long-term	Groundwater	Report status						status	Change Dec to May
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	variation increases with low levels recorded down to c. 6m OD	,								
G3	Located near the base of the cliff. Tip at 13.6m OD. Complex pattern comprising c. 7 month period cycle of rising water level with superimposed sub-weekly fluctuations. 7 month cycle shows rise in water levels of c 1m from 13.3m OD in Oct 2012 to high of 14.4m OD in Feb 2013, falling to low of 13.5m OD in June 2013.	13.2m OD 14.4m OD 1.2m							Groundwater level continues to fall at 13.2m OD in August 2016, before rising and gradually falling to 13.3m OD by October 2016.	Groundwater level increased slightly by March 2017 and remains steady at an average of 13.4m OD.
5 spa	Located near the base of the cliff. Tip at 9.4m OD. No correlation with the upper tip in this well. Data only recorded between Sep 2006 and May 2012, after which the hole is dry. Limited fluctuation between c. 8.5m and c.9.5m OD.	8.5m OD 9.6m OD 1.1m							Borehole dry. Check piezometer integrity.	Borehole dry. Check piezometer integrity.
BH1a spa	Located at the toe of the Spa landslide. Tip at 2m OD. Subweekly fluctuation about mean around 4.4m. Water levels were at their highest during Jan and Feb 2012 when they were c. 0.5m higher than average. Sub-	3.9m OD 5.0m OD 1.1m							Continuing cyclical pattern overlain onto steady groundwater levels during the monitoring period. Fluctuations within range of previous records. Average groundwater level c. 4.5m OD.	Continuing cyclical pattern overlain onto steady groundwater levels during the monitoring period. Fluctuations within range of previous records, greatest between January and March 2017 at 0.5m. Average groundwater level c. 4.5m OD.

	Long-term	Groundwater	Report status						Change Dec to May	
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	weekly fluctuations are c. 0.4m in the period Oct 2012 to Mar 2013.	,,								
BH1b spa	Located at the toe of the Spa landslide. Tip at 10.1m OD. Similar pattern to BH1a. Subweekly fluctuation in water level about mean of c. 12.4m OD. Water levels highest in late Feb 2012 when they reached 12.7m OD. Subweekly fluctuations were up to 0.5m in the period Oct 2012 to Mar 2013.	12.0m OD 12.8m OD 0.8m							Continuing cyclical pattern in groundwater levels. Average groundwater level c. 12.5m OD.	Continuing cyclical pattern in groundwater levels. Reaches new historical high in December 2016 at 12.8m OD, and fluctuates between January and March up to 0.7m. Groundwater levels averaging 12.6 in May 2017. Nearby inclinometer BH101 shows no significant ground movement.
BH1 Prom	Located inland of the cliff top. Tip at 41.4m OD. 5 month period where water-level rose c. 1m from 41.5m OD in Oct 2012 to 42.6m OD in late Feb 2013, followed by period of gradual fall to 41.8 in late 2013. Superimposed on this trend are sub-weekly fluctuations of c. 0.3m.	41.2m OD 43.7m OD 1.4m							Groundwater levels fall rapidly to 42.9m OD in August 2016 and continue to gradually fall in October to 42.7m OD. Groundwater levels remain elevated.	Groundwater level shows a net decrease during the monitoring period. A significant drop in groundwater of c.0.5m levels occurs in February. This is assumed to be a systematic error. Between February and May water levels increase slightly by 0.2. Water levels decrease gradually into June, averaging 42.1m OD, within the historical range. Check piezometer integrity
G1a	Located inland of the cliff top. Dipped piezometer that shows consistent water levels of c. 53.5m OD since late 1997.	53.4m OD 53.9m OD 0.3m							Groundwater level at historical low of 53.4m OD in October 2016.	Borehole dry. Check piezometer integrity
G1b	Located inland of the cliff top. Dipped piezometer that	19.2m OD 51.1m OD 31.9m							Consistent water level since previous	Borehole dry. Check piezometer integrity

	Long-term	Groundwater					Rep	ort	status	Change Dec to May
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	shows significant variability from late 1997 to early 2003 when water levels dropped from c 50m OD to c. 20m OD with significant fluctuations, and subsequent period of consistent level at c. 19m OD. There was a short lived rise to c. 21m during Dec 2012.	Williy Waxy Kange							recording at 19.2m OD.	
BH108a	Deep piezometer tip located midslope. Solinst data logger. Record begins on 18 Dec 2012 and shows several sharp fluctuations that may be a response to rainfall events or ingress of surface water.	9.3m OD 31.4m OD 22.1m							Groundwater levels continue to fall over monitoring period to 23.7m OD by October 2016.	Groundwater levels continue to fall over monitoring period to 23.1m OD by May 2017.
BH108b	Shallow piezometer tip co-located with deeper BH108a. Dry between Sept 2012 and Jan 2013.	25.6m OD 31.6m OD 6m							Water level falls to 29.2m OD, remaining close to historical high.	Borehole dry. Check piezometer integrity
BH106a	Located at the cliff top. Solinst data logger. Borehole dry between Oct 2012 and Jan 2013.								Borehole dry. Check piezometer integrity	Borehole dry. Check piezometer integrity
BH106b	Located at the cliff top. Located at the cliff top. Borehole dry between Oct 2012 and Jan 2013.	n/a (dry)							Borehole dry. Check piezometer integrity	Borehole dry. Check piezometer integrity
BH104a	Located near the base of the slope.	5.0m OD 20.0m OD							Groundwater levels remain steady at	Groundwater levels remain steady at 4.3m OD. No fluctuation.

	Long-term	Groundwater		Report status						Change Dec to May
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	Solinst data logger.	15.0m							4.3m OD. No fluctuation.	
BH104b	Located near the base of the slope. Manual piezometer tube. Borehole dry between Sept 2012 and Jan 2013.	4.3m OD 11.0m OD 6.6m OD							Groundwater level falls to 10.3m OD. Remains close to historical high.	Groundwater level remains steady at 10.3m OD, which is close to historical high.
BH102a	Located at the base of the slope behind the seawall. Solinst data logger.	0m OD 2.6m OD 2.6m							Decrease in magnitude of cyclical change to 0.9m. Average groundwater levels steady at 1.4m OD over monitoring period.	Slight increase in average groundwater levels to 1.6m OD over monitoring period.
BH102b	Located at the base of the slope behind the seawall. Manual piezometer.	1.1m OD 2.1m OD 1.0m							Slight increase in groundwater to 1.9m OD.	Decrease in groundwater to 1.6m OD.

These data indicate:

- Inclinometers show no significant movement in the monitoring period.
- Most locations show continuation of past patterns or slight decreases in water level over the monitoring period, with the exception of BH104b which indicates continued elevated groundwater levels. This trend should be reviewed in the next report.
- Piezometers 5 Spa, G1a, G1b, BH106a, BH106b and BH108b should be checked because they
 were dry. This equipment may be damaged and required attention to determine whether they
 can be repaired.
- Piezometer data from BH108a shows that the earlier trend in rising groundwater level has changed following installation of the new telemetry logger. This shows groundwater is falling well below earlier levels. This trend should be reviewed in the next report.
- Piezometer data from H5 are now collected via telemetry, and show water levels follow the same saw-tooth pattern recorded previously. Sharp increases in groundwater level tend to occur in response to peak rainfall events. An erroneous peak in groundwater level occurs in February. This location should be checked and the trend should be reviewed in the next report.
- Piezometer data from BH1 Prom shows a systematic fall in water levels by 0.5m to a new base level in February. This is likely to be a systemic error and should be checked during the next monitoring period.
- Acoustic emissions (AE) detected are negligible and do not suggest slope movement.
 Fluctuations in the data represent the specific hydrogeological conditions at this site (i.e. time dependent seepage conditions following rainfall events). This is supported by the inclinometer measurements at BH105, which show negligible movement. (Figure 9.2 and 9.3)

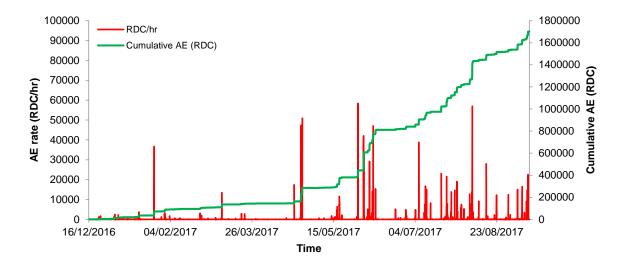


Figure 9.2. Cumulative AE (RDC) and AE rate (RDC/hr) time series at Scarborough Spa for the period December 2016 to September 2017.

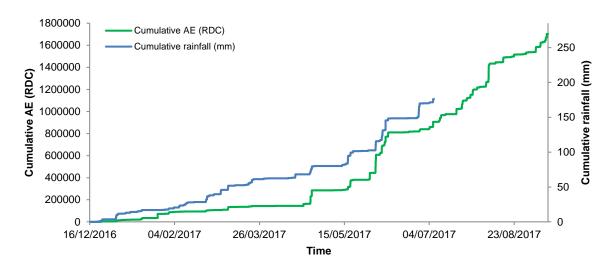


Figure 9.3. Cumulative AE (RDC) and cumulative rainfall time series at Scarborough Spa for the period December 2016 to September 2017.

9.4.4 Clock Café (MU 22/4)

Monitoring at the Clock Café comprises a line of three boreholes from the promenade (BH15) to the midslope (AA10 F2) and lower slope (AA11 F4) (Table 9.8, Figure 9.1B).

Table 9.8. Summary of inclinometer data at the Clock Café

Barrelada	Comment of the Adams			R	eport	statu	s	Change Dec to May 2017
Borehole	Summary of past data	1 2 3 4 5 6 7		Change Dec to May 2017				
AA10 (F2)	30m deep borehole through 3m of made ground, 21m of glacial sediment and 6m of siltstone/sandstone bedrock at the headscarp of the Clock Café landslide. Ground level is 34.98m OD, base of hole is 4.98m OD. Very low creep indicated in the upper 5m, with incremental displacements of up to 5mm. 30 June 2012 reading is erroneous.						No significant change.	No data. Collect data on next site visit.
AA11 (F4)	20m deep borehole penetrating 8m of glacial sediment and 12m of siltstone/sandstone bedrock near the toe of the Clock Café landslide. Very low cumulative movement along whole length of borehole of up to 3mm is within tolerance of the device.						No significant change.	No significant change.

Table 9.9. Summary of groundwater data at the Clock Café

	Long-term	Groundwater				R	eport	t Statı	ıs	Change Dec to May
Borehole	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
BH15	Located inland of the landslide headscarp. No historical data	n/a							Borehole dry. Check piezometer integrity	Borehole dry. Check piezometer integrity

The data show no ground movements at the Clock Café, which is a continuation of the past pattern of stability at this location. Data from inclinometer AA10 is to be collected on the next site visit. The one piezometer at this location continues to be dry. This equipment may be damaged and required attention to determine whether it can be repaired.

9.4.5 South Cliff Gardens (MU 22/5 and 22/6)

The South Cliff Gardens area comprises landscaped public areas and the former South Bay Pool, which lies at the foot of a relict landslide complex (the South Bay Pool landslide). There are three transects of monitoring locations (Tables 9.10 and 9.11; Figure 9.1C).

Table 9.10. Summary of inclinometer data at South Cliff Gardens

						Report	statu	s	
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017
AA08 (D3)	24m deep borehole that penetrates 12m of glacial sediment and 12m of interbedded bedrock. Ground level is 38.43m OD, base of hole is at 14.43m OD. Data indicate slight progressive creep along the whole length of the borehole,							No significant change.	No significant change.
BH17	50m deep borehole than penetrates 34m of glacial sediment and 16m of siltstone bedrock at the top of a mudslide embayment. Ground level is 57.46m OD, base of hole at 7.46m OD.							No significant change.	No significant change.
BH16A	54m deep borehole than penetrates of 33m of glacial sediment and 21m of siltstone/sandstone bedrock inland of the Rose Garden rotational landslide. Ground level is 62.88m OD, base of hole is 8.88m OD.							No significant change.	No significant change.
вн20	41m deep borehole that penetrates 27m of glacial sediments and 14m of sandstone bedrock within the body of a small landslide. Ground level is 58.98m OD, base of borehole is 17.98m OD.							No significant change.	No significant change.

Table 9.11. Summary of groundwater data at the South Cliff Gardens

		Groundwater					Change Dec to May			
Borehole	Long-term Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
BH18a	Tip at 26.8m OD near the base of the cliff and Rose Garden landslide. Complex pattern, with sub-weekly peaks 4m to 5m higher than base readings associated with storms. From	34.4m OD 42.5m OD 6.1m							Continuation of past pattern of short lived spikes. In August 2016 a spike ranged 3.7m. Base levels have gradually decreased over the monitoring period to 34.5m OD in October close to the historical	Groundwater rises rapidly from an average of 34.5m OD in Dec to c. 39m OD in late February to March 2017, where it stabilises in an elevated position relative to earlier records.
	Nov 2012 to May 2013 base readings were 37m OD. Between May and								low.	

		Groundwater					Rep	oort St	atus	Change Dec to May
Borehole	Long-term Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	Aug 2013 levels rose to 38m OD.	·····y ·····y								
BH18b	Tip at 23.8m OD near the base of the cliff and Rose Garden landslide. Pattern very similar to BH18a installed higher in the borehole	34.3m OD 42.4m OD 6.1m							Same pattern and water levels as BH18a. Spikes on same days, suggesting connectivity between both piezometer tips and possible ingress of surface water to the borehole. Base levels falling over monitoring period to 34.3m OD. Recommend integrity of installation is checked.	Previous records show the same pattern and water levels as BH18a. These now differ over this monitoring period whereby there is only a slight rise in average groundwater levels to 34.8m OD, close to the historical low.
BH19a	Tip at 53.8m OD inland of the headscarp of the South Bay Pool landslide. This piezometer has been dry since installation.	52.8m OD 53.3m OD 0.5m OD							Cyclical pattern with magnitude of variation ranging 0.4m averaging 53.1m OD. Peaks in groundwater levels in August and October 2016.	Cyclical pattern with magnitude of variation ranging 0.7m averaging 52.9m OD. Peaks in groundwater levels in December 2016 and falls in February.
BH19b	Tip at 47.3m OD inland of the headscarp of the South Bay Pool landslide. Submetre variation about an average level of 47.8 OD. Periods of slightly higher water level from Dec 2012 to Mar 2013, late May 2013 and early Aug 2013.	47.1m OD 49.3m OD 2.2m							Continuation of sub- weekly fluctuation of c. 1.m about monthly variation. Levels spike to historical high in late June at 49.3m OD and in August 2016. Overall decline in base levels, minimum at 47.2m OD.	Continuation of sub- weekly fluctuation of c. 1.m about monthly variation. Overall steady decline, averaging 47.5m OD.
D2a	Tip at 27.5m OD at the headscarp of the South Bay Pool landslide. Submetre variation about an average level of 40.5m OD. Periods where hole appears dry occurred regularly from late June to early July 2013, following which no data has been recorded.	31.1m OD 40.9m OD 9.8m							No data since May 2016, logger to be checked.	Levels remain steady around 31.6m OD with sub-weekly fluctuations of up to 0.5m.

Danahala	Lawa tawa Battawa	Groundwater					Rep	ort St	atus	Change Dec to May
Borehole	Long-term Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
D2b	Tip at 41.5m OD at the headscarp of the South Bay Pool landslide. Pattern similar to that recorded by lower elevation tip, with sub-metre variation about mean of c. 45.8m OD. Slight peak in water level occurred in late Nov to late Dec 2012. Gap in data between April and Aug 2013.								No data, logger to be checked.	No data, logger to be checked.
Bh3a	Tip at 41.5m OD at a mid-slope position adjacent to the South Bay Pool landslide. Submetre variation about a mean value. Change occurs in Apr 2013, before which mean is 44.5m OD, after which it is drops to c. 44m AOD.	Original logger: 43.6 44.8 1.2 Replacement logger: 45.6m OD 50.0m OD 4.4m							No data, logger to be checked.	No data, logger to be checked.
Bh3b	Tip at 10.5m OD at a mid-slope position adjacent to the South Bay Pool landslide. Similar pattern to high elevation tip, however uniform level of 10.5m OD is interrupted by frequent short-duration (1 day) peaks that are up to 8m higher. Peaks particularly common during period Nov 2012 to Feb 2013 and May to June 2013.	10.3m OD 18.6m OD 8.3m							No data since May 2016, logger to be checked.	No data since May 2016, logger to be checked.
E2a	Tip at 31.4m OD below the headscarp of the mudslide embayment. Cyclical long-term pattern with sub-	43.3m OD 46.5m OD 3.2m							Groundwater level gradually falls to 44m OD throughout monitoring period.	Groundwater level gradually continues to fall to 43.5m OD in March, where it stabilises at 43.7 in May 2017.

		Groundwater					Rep	ort St	atus	Change Dec to May
Borehole	Long-term Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
	metre fluctuations superimposed. Water levels rise from c. 44m AOD to 46.5m OD between Oct 2012 and late Feb 2013 thereafter they fall gradually to 44.7m OD in Oct 2013									
E2b	Tip at 43.6m OD below the headscarp of the mudslide embayment. Different pattern to shallower tip, with sub-metre variation about a mean of 51m OD.	49.6m OD 51.4m OD 1.7m							Groundwater levels remain steady averaging c. 50.4m OD.	Groundwater levels remain steady averaging c. 50.4m OD, with the exception of a rapid fall in March 2017 by 0.5m.

These data indicate:

- No movement has been recorded in any boreholes at South Cliff Gardens.
- Water levels are generally stable or show gradual decreases, reflecting the antecedent rainfall occurring during in the winter and spring.
- Shallow piezometer BH18a shows a rapid increase in groundwater levels relative to deeper BH18b which shows a more gradual increase during February 2017. This suggests surface water ingress and the contractor should ensure that plastic caps are in place and that water cannot collect at the top of the boreholes. These piezometer data should be reviewed during the next monitoring period. There are no significant ground movement in nearby boreholes.
- Borehole piezometers D2a, D2b, BH3a and BH3b should be checked and repaired.

9.4.6 Holbeck Gardens (MU 22/7)

This area comprises two monitoring locations (Figure 9.1C); water levels are monitored at two depths along the promenade and ground movements are recorded by an inclinometer on the upper slope (Tables 9.12 and 9.13).

Table 9.12. Summary of inclinometer data at Holbeck Gardens

Borehole	Summary of past data					Repo	rt stat	us	Characa Dan ta Mara 2017
		1	2	3	4	5	6	7	Change Dec to May 2017
AA07 (BH2)	60m deep borehole penetrating 31m of glacial sediments and 29m of siltstone/sandstone bedrock. Ground level is 56.33m OD, base of hole is -3.67m OD. Data show progressive displacement of the glacial sediments, with up to c. 15mm at the ground surface. There is a suggestion of a shear developing at the contact between the glacial sediments and underlying bedrock and also at c.14m depth, within the glacial sediments.							No significant movement.	No significant movement.

Table 9.13. Summary of groundwater data at Holbeck Gardens

Borehole	Long-term	Groundwater				R	epor	t stat	us	Change Dec to May
	Pattern	summary Min/Max/Range	1	2	3	4	5	6	7	2017
Bh4a	Tip at 31.5m OD. Complex pattern with periods of stability interspersed by rapid rises or falls of up to 2m.	47.1m OD 58.8m OD 11.7m							Unable to download logger data. Check integrity of piezometer.	Unable to download logger data. Check integrity of piezometer.
Bh4b	Tip at 35m OD. Different pattern to records of shallower tip. Highly variable	31.8m OD 59.9m OD 26.7m							Groundwater levels fall to 57.7m OD in August 2016 before rising to 58.4m OD in September. Levels return to 57.7m OD at end of monitoring period. Levels remain elevated.	Unable to download logger data. Check integrity of piezometer.

The data logger was at fault for Bh4a and Bh4b, and data were not downloaded. The integrity of the piezometers should be checked. No evidence of movement is shown in the current inclinometer data.

9.5 Causal-response relationships

Groundwater levels tend to show a slight decrease or remain steady during the monitoring period, which reflects the relatively mild and dry conditions from December 2016 onwards. Groundwater level generally peaks in December, following the antecedent heavy rainfall during November 2016.

There is little evidence of movement in the inclinometers and no critical groundwater level thresholds have been identified during this period.

9.6 Implications and recommendations

None of the inclinometers indicate ground movement.

The majority of piezometers show groundwater levels have either been maintained or decreased. Several show short-lived peaks in water level that suggests ingress of surface water during heavy rainfall. Checks should be made at these locations to ensure water-proof caps are in place. At South Bay Gardens in borehole BH18a, groundwater levels have risen to an elevated position relative to earlier records, however no significant ground movement was recorded by nearby inclinometers. This record should be reviewed during the next monitoring period to check whether this trend continues. Similarly, elevated groundwater levels in BH104b at Scarborough Spa should be checked and reviewed during the next monitoring period.

No data were collected at a number of piezometers including BH12 (Spa Chalet), D2b, Bh3a, Bh3b, Bh4a and Bh4b (Clock Café). Boreholes 5 spa, G1a, G1, BH108b. BH106a, BH106b (Spa) and BH15 (Clock Café) are recorded as dry. The integrity of piezometer tips should be checked and the next monitoring data reviewed, whether these trends continue. Data for piezometers H5 and BH1 Prom at the Spa are also unavailable for this monitoring period.

Inclinometer data at borehole BH13 at the Spa appears erroneous. It is likely that the borehole is blocked and requires clearing. No data is available for inclinometer in borehole AA10 at the Clock Café and data should be collected during the next site visit.

Filey Town

10.1 Site description

The cliffs at Filey are formed in thick (c. 50m) glacial sediments that overlie the Upper Jurassic Kimmeridge Clay Formation across the town frontage and Upper Calcareous Grit north of the town towards Filey Brigg. The cliffs are protected by a sea wall at Filey and unprotected to the north and south of the town. Outflanking of the seawall and cliff instability of both the protected and unprotected cliffs at Filey is a concern. The cliffs across the town frontage have been landscaped and are criss-crossed with public footpaths. The glacial sediments have been deeply incised to form Church Ravine to the north of the town and Martin's Ravine to the south.

In July 2007, an intense rainstorm resulted in severe and widespread flooding throughout Filey; the storm water run-off caused many slope failures and extensive scour damage to paths and bridge abutments within Martin's Ravine. Existing drainage was overwhelmed and extensively damaged due to the excessive storm water run-off around Glen Gardens and this also caused drainage to collapse leading to slope instability behind Royal Parade chalets and Crescent Hill (Mouchel, 2012). The unprotected cliffs to the north and the south of the town are susceptible to toe erosion by the sea and are actively retreating. Cliff behaviour units (CBUs) have been defined and their activity status classified under the Cell 1 Regional Monitoring Programme.

10.2 Ground model and monitoring regime

Cliff behaviour units, reflecting individual mudslides and areas of relict cliff protected by the seawall, have been mapped for the frontage (Figure 10.1):

- MU29/AA and /AB are cliffs and mudslides south of the town
- MU 28/Z is a till cliff protected by rock armour immediately south of the sea wall
- MU27/X and MU28/Y are dormant cliffs protected by the seawall
- MU27/T /U, /V and /W are cliffs and mudslides north of the town

Halcrow (2012a) provides an overview of the ground models throughout the Filey Town frontage. The whole cliff line is comprised of weak glacial sediments that tend to fail through simple landslides triggered by both toe erosion and elevated groundwater levels.

The cliffs at Filey town, which are protected by a seawall, display evidence of historical instability. Shallow failures last occurred in this area in response to the intense storm event of July 2007.

Within the ravines, the steep till slopes are susceptible to shallow failure resulting from toe undercutting and excess groundwater levels due to intense and prolonged rainfall events.

The monitoring regime at Filey Town comprises the following:

- Filey Park Till cliff with ground water monitoring at the cliff top.
- Golf Course Ground water monitoring at the cliff top.
- Church Ravine/Coble Landing Ground water monitoring at the cliff top and an inclinometer at the cliff toe.
- The Crescent/Rutland St Groundwater monitoring at the cliff top and an inclinometer at the cliff toe.
- Glen Gardens/Martin's Ravine Ground water monitoring on the cliff top and toe. Inclinometers at the cliff top and toe.

- Muston Sands Ground water monitoring at the cliff top.
- Inland North Groundwater monitoring near Church Cliff Farm, Pinewood Avenue and Parish Wood.
- Inland South Groundwater monitoring near Filey Fields Farm, Long Plantation (west of Rivelin Way and Fewston Close) and Filey School.

10.3 Historical ground behaviour

Filey town was monitored by Mouchel Ltd for the period between summer 2009 and summer 2012. A summary of their results is provided in Table 10.1, which shows minor movement in one borehole during the autumn of 2009 but without subsequent movement and limited fluctuation in ground water level which Mouchel attribute to tidal variation in some boreholes and variations in stream flow in others. No relationship between groundwater level and ground movement was reported by Mouchel. Additional monitoring covering the period April 2011 to Dec 2012, associated with the recent seawall outflanking study, are provided in Halcrow (2013a).

Table 10.1 Summary of historical ground behaviour at Filey Town.

Observations in Mouchel 2012 (covering 6 month period Total Change observed between July 2009 and June 2012 between Dec 2011 and June 2012 Groundwater levels in BH5B (toe of Glen Gardens/Martin's Mouchel report that ground water levels have increased Ravine) and BH6 (midslope Glen Gardens/Martin's Ravine) since December 2011, the maximum rise having been rose by 49mm and 560mm respectively. BH1 (cliff top identified as 560mm at BH4, Mouchel also describe erratic Glen Gardens/Martin's Ravine, now inactive) rose 152mm readings from this borehole. Mouchel describe an which appeared to reflect prevailing water level in increase of 49mm at BH5b and attribute this to tidal Martin's Ravine. BH04 (midslope Glen Gardens) was noted fluctuations. Ground water readings from BH1 and BH2 to be recording erratically. The inclinometer in BH3 was appear to have remained relatively constant at about 15m not readable during this time and no new movement was OD. Only 'baseline' inclinometer readings have been reported at BH6. determinable from BH3. Mouchel observe that ground water readings from BH1 seem to reflect water levels within the stream flowing in Martin's Ravine. Initially (between September and December 2009), displacements of <5mm were noted in BH6 but no further movements have been identified.

10.4 New data

Tables 10.2 and 10.3 summarise the inclinometer and piezometer data from Filey Town to June 2017.

These data indicate:

- No movement has been recorded in any boreholes at Filey Town.
- Water levels are generally stable or decreasing slightly in most boreholes.
- Groundwater levels have risen again in borehole CPBH01a by 0.9m near to the historical high, following a previous fall to lower levels. It is possible the borehole was flooded by surface water flows. Groundwater levels also remain elevated in boreholes CPBH02a, CPBH09a and CPBH10a.
- Dip meter reading at CPBH09b did not reflect water levels recorded by the diver, suggesting a systematic error. The spreadsheet has been amended to correct water levels.
- Boreholes CPBH10b and CPBH08b were dry.
- CPBH01b (Diver), CPBH04 (Diver), CPBH06b (Diver) and CPBH03 (inclinometer) were not downloaded for this monitoring period.

Note boreholes BHA, BHB, BHC, BHD, TP3, TP6 and TP9, which are inland of the coast and have a focus on flood risk, are no longer included in the coastal instability monitoring programme.

Table 10.2. Summary of inclinometer data at Filey Town.

					F	Repor	t sta	tus	Charge Doots May 2017	
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017	
СРВН03	CPBH03 is 10m deep. Surface elevation is c. 6m OD* therefore the base of the borehole is at -4.0m OD* and extends through 4.4m of made ground and 5.6m of glacial sediment.							No significant movement.	No data, borehole obstructed. Collect on next site visit.	
СРВН05	CPBH05 is 10m deep. Surface elevation is c.6.5m OD* therefore the borehole extends to ca3.5m OD* through glacial sediments.							No significant movement.	No significant movement.	
RCBH07	CPBH07 is 20m deep. Surface elevation is at c. 5m OD* through glacial sediments.							No significant movement.	No significant movement.	
вн6	BH6 is 30m deep. Surface elevation is c.27.4m OD* therefore the base of the hole is c2.6m OD. The borehole extends through glacial sediment. Cumulative displacements of 10mm in a negative B axis between Sept and Dec 2009 likely to be error.							No significant movement.	No significant movement.	

Note: *Surface elevation and borehole depth calculated from digital elevation model.

Table 10.3. Summary of groundwater data at Filey Town

Borehole	Long-term Pattern	Groundwater summary				F	Repor	t Sta	itus	Change Dec to
		Min/Max/Range	1	2	3	4	5	6	7	May 2017
вн5b	Tip depth at 1.09m OD. Levels constant with limited fluctuation between 1.1m OD (Aug 2008) and 1.7m (Dec 2009).	1.1m OD 7.5m OD 6.4m							Levels increase slightly to 1.6m OD during monitoring period.	Groundwater levels fall slightly to 1.5m OD by May 2017.

Borehole	Long-term Pattern	Groundwater summary				itus	Change Dec to			
		Min/Max/Range	1	2	3	4	5	6	7	May 2017
вн4	Tip at 18.07m OD. Major fluctuations (>7m) in groundwater elevation between Dec 2009 and June 2011. Mouchel (2012) have previously reported groundwater readings from this piezometer as 'erratic'. Readings have been more settled 2011 albeit showing an increase to 20.2m OD in May 2012.	19.7m OD 27.1m OD 7.4m							Levels slightly decrease to 21m OD.	Groundwater levels decrease to 20.6m OD by May 2017.
СРВН01а	Readings sporadic BH often dry. Mean level is 17.17m OD, with variation between 16.89m OD (15/12/2011) and 17.48m OD (20/12/2012). This latter measurement is likely to reflect the cumulative impact of the wet spring, summer and winter of 2012.	16.9m OD 26.2m OD 8.3m							Groundwater levels fall near to historical low of 17.1m OD.	Groundwater level rises by 0.9m close to historical high at 26.1m OD.
CPBH01b (Diver)	Tip at 32.63m OD. Fluctuating but steadily rising water level from 33m OD in late 2011 to 34m OD in summer 2012. Slight drop in autumn 2012 before sudden rise to maximum of 35.0m OD on 14 Dec 2012.	33.0m OD 35.0m OD 2.0m							No data. Readings to be re-taken. Manual dip readings at 33.4m OD, which suggests levels are near the historical low.	No data. Readings to be re-taken. Manual dip readings at 33.9m OD, which suggests levels remain within the range of previous records.
СРВН02а	Tip at 1.57m OD. Mean groundwater elevation at c. 5m OD with minor fluctuations. Short lived drop to 3.57m in Sept 2012. Maximum level 5.23m OD on 19/04/2012.	3.6m OD 5.2m OD 1.6m							Water levels increase slightly to 5m OD, close to historical high.	Groundwater levels fall slightly to 4.9m OD, remaining at an elevated position.

Borehole	Long-term Pattern	Groundwater summary				ı	Repor	t Sta	itus	Change Dec to
		Min/Max/Range	1	2	3	4	5	6	7	May 2017
CPBH02b (Diver)	Tip at 8.17m OD. Generally steady at c. 8.7m OD except for spikes in on 6 July 2012 (to 15.6m OD) and 7 Dec 2012 (to 20.0m OD). Smaller spikes (to c. 9.7m OD in late Nov/early Dec 2012).	5.1m OD 20.0m OD 14.9m							Water level steady at 8.6m OD.	Water level steady at 8.6m OD, which is credible and reflects past records, but manual dip readings indicate borehole is dry. Recommend installation integrity is checked.
СРВН04а	Tip at 2.90m OD. Mean ground water level at 7.2m OD, with range of fluctuation between 7.02m OD (06/09/2012) and 7.33m OD (19/04/2012).	7.1m OD 32.9m OD 25.8m							Water levels remain steady at 7.2m OD.	Groundwater levels fall slightly to 7.1m OD.
CPBH04 (Diver)	Tip at 9.9m OD. Steady around 13.5m OD until Dec 2012 although dip in Dec2012 reads significantly higher (16.3m OD).	13.2m OD 13.6m OD 0.4m							No data. Readings to be re-taken. Manual dip readings at 13.3m OD, which indicates levels are static.	No data. Readings to be re-taken. Manual dip readings at 13.3m OD, which suggests levels are static.
СРВНО6а	Tip depth at 0.13m OD. Mean groundwater elevation at 19.86m OD. Range between 18.85m OD (27/02/12) and 20.11 (20/12/12). Notable increase in Mar/April 2012 followed the dry period of late autumn 2011 to winter of 2011/12. Rises to highest level in Dec 2012 after very wet year.	18.9m OD 20.0m OD 1.1m							Water level falls slightly to 18.9m OD.	No data, readings to be re-taken. Manual dip readings indicate groundwater levels rise slightly to 19.1m OD.
CPBH06b (Diver)	Tip depth at 8.63m OD. Steady at c. 18m OD except for sudden drop to around 14.5m OD and immediate recovery on 20/03/12 and 06/09/12 and sudden drop on 19/04/12 followed by a prolonged steady period at c. 15m OD before sudden recovery on 24/05/12 to 18m OD.	9.2m OD 19.1m OD 9.9m							No data. Readings to be re-taken. Manual dip reading at 18.9m OD, which indicates continuation of historical water level.	No data. Readings to be re-taken.

Borehole	Long-term Pattern	Groundwater				ı	Repor	t Sta	tus	Change Dec to
		summary Min/Max/Range	1	2	3	4	5	6	7	May 2017
СРВН08а	Mean groundwater elevation is 8.71m OD ranging between 8.48m OD (19/04/2012) and 9.46m OD (20/12/2012), suggesting a greater lag time or less responsiveness to antecedent rainfall conditions.	8.5m OD 11.4m OD 1.9m							Slight increase in water level to 9.6m OD.	Groundwater levels remain steady at 9.6m OD.
CPBH08b (Diver)	Very steady with fluctuations over whole period only between 17.90m OD and 17.97m OD.	17.7m OD 19.3m OD 1.6m							Water levels static at 17.9m.	No data. Readings to be re-taken. Dip meter data indicates borehole is dry. Recommend installation integrity is checked.
СРВН09а	Tip depth at 0.64m OD. Mean groundwater elevation is 20.27m OD and ranges between 19.86m OD (01/08/2012) and 20.98m OD (06/09/2012).	19.9m OD 21.0m OD 1.1m							No change in water-level which is at 20.2m OD. Dip meter readings corroborate diver data.	Groundwater levels rise slightly to 20.4m OD.
CPBH09b (Diver)	Tip Depth at 17.74m OD. Between 01/01/2012 and 20/12/2012 levels fluctuate between 19.9m OD and 20.5m OD. There is a general trend of slight decline towards June 2012 followed by a rise towards peaks in late Oct and mid- Dec 2012.	18.8m OD 21.1m OD 2.3m							Spreadsheet error detected and resolved. Water-levels static at c. 20.5m, which is corroborated by dip meter data.	Groundwater levels fall to 20.1m OD, which is corroborated by dip meter data.
CPBH10a (Diver)	Tip depth at 23.82m OD. Shows pattern of sharp increases over a week, followed by gentle decreases over several weeks, to c. 28.5m OD. Comparison to rainfall records indicates borehole has a comparatively 'flashy' response to rainfall, with lag times reducing towards the end of 2012 as atypically low groundwater levels recovered. Max peak is 30.8m OD in late Dec 2012.	24.6m OD 30.8m OD 6.2m							Continuation of saw-tooth pattern, generally decreasing groundwater levels. Slight increase in August to 28.7m OD. Decreasing for remainder for monitoring period to 28.5m OD.	Continuation of saw-tooth pattern, with gradual increase in underlying groundwater level. Peaks in March 2017 at c. 29.4m OD.

Borehole	Long-term Pattern	Groundwater summary				tus	Change Dec to				
		Min/Max/Range	1	2	3	4	5	6	7	May 2017	
СРВН10Ь	Tip depth at 11.92m. No data prior to October 2013 due to blockage by slip rod.	n/a (dry)							Borehole dry Recommend installation integrity is checked.	Recommend installation integrity is checked.	

10.5 Causal-response relationships

Most piezometers show a weak response to rainfall with the exception of shallow piezometers CPBH01a and CPBH10a which respond rapidly within a month to peaks in rainfall. Muted antecedent rainfall responses are also noted in CPBH06a and CPBH09a. Most of the piezometers show steady groundwater levels or slight decreases, reflecting the relatively dry and mild conditions during the monitoring period. There has not been movement in inclinometers and therefore no relationships between groundwater and ground movement have been identified.

10.6 Implications and recommendations

No data are available for boreholes CPBH01b, CPBH04 CPBH06a and CPBH06b and require readings to be re-taken on the next site visit. Boreholes CPBH02b, CPBH08b and CPBH10b were dry and require investigation and repair work where necessary. The rising trend in groundwater levels towards the historical high in CPBH01a and CHBH10a should be monitored and reviewed in the next monitoring report. Groundwater levels in CPBH02a and CPBH09a remain elevated and should also be reviewed in the next report to see whether this trend continues.

Inclinometer at BH6 appears to have debris at its base, causing error at the lower part of the reading. Borehole CPBH03 was obstructed during the site visit and data should be collected for the next monitoring report.

Filey Flat Cliffs

11.1 Site description

Flat Cliffs is a private residential settlement located on coastal slopes in central Filey Bay. The settlement includes private homes and a Yorkshire Water pumping station accessed via a private road down the cliffs that is particularly steep near the top of the cliffs (Halcrow, 2012b). The cliffs are formed in thick and variable glacial sediments that continue to at least 12.4m below OD and which are prone to cliff instability. There is concern that ongoing cliff instability threatens properties and the only access road to about 40 homes at Flat Cliffs (Halcrow, 2012b).

11.2 Ground model and monitoring regime

This site comprises three cliff behaviour units: MU29/AQ, which is an active mudslide complex north of the main settlement and MU29/AR and MU29/AS that form the main landslide undercliff upon which the settlement has been developed.

The undercliff ground model can be described as a complex landslide system that is backed by a steep headscarp and fronted by a sea-cliff (Halcrow, 2012b). The undercliff morphology comprises landslide scarps and benches, some of which are back-tilted although interpreted as failing on translational shear surfaces rather than rotational failure. A large mudslide complex in the north of the site is periodically active, and threatens the access road and properties. Activity is generally associated with accelerated toe erosion and elevated groundwater levels.

The monitoring regime at Flat Cliffs includes the following (Figure 11.1):

- North of site automated piezometer on the cliff top and inclinometer on the access road.
- Central site Piezometers with data loggers on the cliff top and next to the access road in the lower slope. Two inclinometers either side of the main access road (Flat Cliffs Road and Lower Flat Cliffs) on the coastal slope (one of which is an experimental acoustic inclinometer installed by Loughborough University).
- South of site Co-located automated piezometer and inclinometer on the Lower Flat Cliffs part of the coastal slope.

11.3 Historical ground behaviour

Filey Flat Cliffs was monitored by Mouchel Ltd for the period between summer 2009 and summer 2012. A summary of their results is provided in Table 11.1, which shows some movement in Borehole A2. No relationship between groundwater level and ground movement was reported by Mouchel. Additional monitoring covering the period April 2011 to Dec 2012, associated with a landslide investigation, are provided in Halcrow (2013b).

Table 11.1. Summary of historical ground behaviour at Flat Cliffs

Observations in Mouchel 2012 (covering 6 month period between Dec 2011 and June 2012)	Total Change observed between July 2009 and June 2012						
Mouchel monitored inclinometer A2 during this period and reported no movement. Mouchel report a groundwater level reading from B1 in June 2012 as revealing a reduction of 520mm relative to December 2011. The report mentions that groundwater readings up to May 2012 are reported in Appendix E to that report, but no readings after June 2010 are identifiable from the graph.	Deviation of 15mm near the surface indicated in A2 between December 2010 and June 2011. This had increased by a further 5mm to 20mm by December 2011. No specific comment is made on ground water levels but it appears from the chart in the appendix that ground water levels remain relatively constant at piezometers A2, A3 and D2, with minor fluctuations in B1 and major fluctuations in D1.						

11.4 New data

Tables 11.2 and 11.3 summarise the monitoring results from inclinometers and piezometers at Flat Cliffs up to May 2016.

Table 11.2. Summary of inclinometer data at Flat Cliffs. *Surface elevations and borehole depths calculated from

digital elevation model.

							Repo	rt status		
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017	
A2	A2 is 27.5m deep (surface elevation at 17.93m OD) and extends through glacial sediment. Moderate movements (<5mm cumulative) between Dec 2009 and Dec 2010 increase by a further c. 10mm by June 2011 at shear at c. 6m to 7m OD							No significant movement.	No significant movement.	
C1	C1 is c. 25m deep. Surface elevation is 25.7m OD* the base of the hole is c. 0.7m OD. Shows very minor (<2mm cumulative) displacements up to and including October 2012.							No significant movement.	No significant movement.	
C2	C2 is c. 21m deep. Surface elevation is at 16.5m* and the borehole extends to - 4.5m OD through glacial sediments. Displacements to Oct 2012 within margin of instrument error							No significant movement. Minor displacement extending to c. 1.5m depth, likely relatively shallow surface creep in glacial sediment.	No significant movement. Minor displacement extending to c. 1.5m depth, likely relatively shallow surface creep in glacial sediment.	
C5	C5 is c. 16m deep. Surface elevation is 12.0m OD* and the borehole extends to - 4.0m OD passing through variable glacial sediments. No movement to Oct 2012 apart very minor displacement in the uppermost 1.5m							No significant movement. Minor displacement extending to c. 1m depth, likely relatively shallow surface creep in glacial sediment.	No significant movement. Minor displacement extending to c. 1m depth, likely relatively shallow surface creep in glacial sediment.	

Barrelanda	6						Repo	rt status	Charres Banks Mars 2017
Borehole	Summary of past data	1	2	3	4	5	6	7	Change Dec to May 2017
C1A	Acoustic inclinometer. The Acoustic Emissions (AE) monitoring has not detected any movement of the landslide to the end of 2012. Higher than average rainfall from April to Dec 2012 had no impact on ground movement. The AE monitoring and inclinometer measurements are consistent							AE measurements between Feb 2016 and Dec 2016 do not show significant slope movements. The period of elevated AE activity after June 2016 could be indicative of minor deformations, but it is likely that this is a function of the deterioration of the cover lid generating spurious AE. The lid was replaced on 11 Aug 2016. The elevated bursts of activity recorded in Nov are thought to be caused by vandalism, with residents reporting youths kicking the equipment.	AE measurements between Dec 2016 and Sept 2017 do not appear to show significant slope movements. The periods of elevated AE activity could be indicative of deformations but of a very slow rate and magnitude. It is more likely that this increased AE activity was caused by extraneous noise (e.g. interference with the surface cover).

Table 11.3. Summary of groundwater data at Flat Cliffs

Borehole	Long-term Pattern	Groundwater				F	epor	t sta	tus	a
		summary Min/Max/Range		2	3	4	5	6	7	Change Dec to May 2017
B1	Tip Depth at - 7.64m OD. Monitored since July 2001. Fluctuates between c. 11.2 m OD and 15.6m OD with peaks in July 2003, April 2004 and Dec 2010. Groundwater at 12.9m OD in May 2012.	11.2m OD 15.6m OD 4.4m							Groundwater level continues to fall to 12.9m OD in October 2016.	Groundwater level at 13.4m OD, well below historical high.
D1	Tip depth at 15.61m OD. Monitored with data loggers since late 2011. Levels show large fluctuations between 15.7 m OD (Sept 2008) and 38.4m OD (Mar 2010). Peaks of 28.2m OD in July 2012 and 24.5m OD in early Jan 2012. Mean base groundwater level is 18 to 18.5m OD.	18.1m OD 29.9m OD 11.8m OD							Groundwater levels fall to 18.6m OD in October 2016 and remain steady.	Groundwater peaks at 21m OD in Jan and fluctuates up to 0.5m around an average of 18.6m OD.

Borehole	Long-term Pattern	Groundwater				F	Repor	t sta	tus	
		summary Min/Max/Range	1	2	3	4	5	6	7	Change Dec to May 2017
АЗ	Tip depth at 6.37m OD. Monitored since March 2001. Dipped readings show static ground water level at c. 18.75m OD with for peaks in July 2001 (19.8m OD) and Dec 2010 (21.4m OD) and a low in July 2008 of 11.63m OD. Vibrating wire piezometer installed in Sept 2011 shows static groundwater level of c. 18.0m OD with minor fluctuation.	17.7m OD 18.2m OD 5.0m							No significant change with continuation of 4 to 6 week fluctuations. Groundwater levels remain steady around c. 18m OD.	No significant change with continuation of 4 to 6 week fluctuations. Groundwater levels remain steady around c. 18m OD.
C4a	Tip depth at -3.7m OD. Monitored since Sept 2011. Levels vary between 7.5m OD and 8.4m OD in response to short and medium term tidal cycles (ca. 6 hourly and 4-weekly).	7.5m OD 8.5m OD 1.0m							No data collected since May 2016. Readings to be re-taken.	No data collected since May 2016. Readings to be re-taken.

The new data indicate:

- No evidence for ground movements is shown by inclinometers.
- AE measurements between December 2016 and September 2017 do not show significant slope movements. The periods of elevated AE activity, which are shown by periods of increased gradient of the cumulative RDC record, could be indicative of deformations but of a very slow rate and small magnitude. However, it is more likely that this increased AE activity was caused by extraneous noise (e.g. interference with the surface cover).
- Groundwater data show no significant change. At borehole D1 near the top of the slope adjacent to the main access road water levels remain steady after decreasing significantly since winter 2015/16.
- No data collected at piezometer in borehole C4a, and requires readings to be retaken.

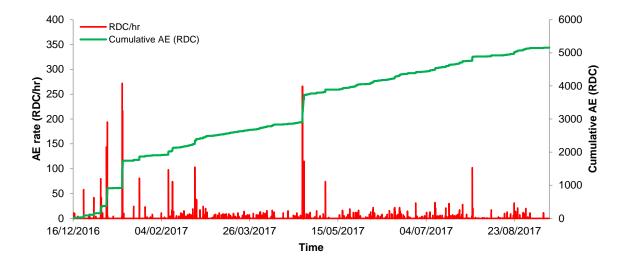


Figure 11.2 Cumulative AE (RDC) and AE rate (RDC/hr) time series at Flat Cliffs, Filey for the December 2016 to September 2017.

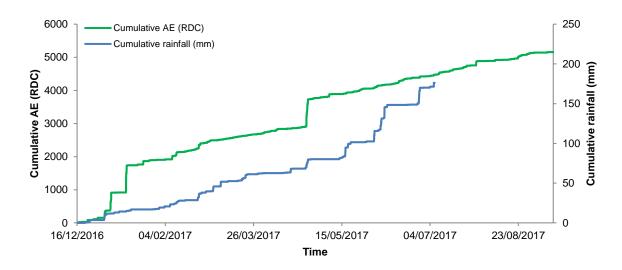


Figure 11.3 Cumulative AE (RDC) and cumulative rainfall time series at Flat Cliffs, Filey for the period December 2016 to September 2017.

11.5 Causal-response relationships

No relationship is identifiable between ground movements and rainfall as no substantial ground movements have occurred. However, borehole D1 appears to show a response to above average rainfall in January and February 2014 and borehole C4a clearly shows the effect of the 5 December 2013 storm surge on groundwater levels as the highest peak in the record. Much of the current monitoring period experienced lower than average rainfall. The data from borehole D1 record a peak in January 2017, which may be in response to over two-month antecedent rainfall. Unusually no peak is recorded associated to August 2016 rainfall. B1 gradual decrease in groundwater level follows a month antecedent rainfall.

11.6 Implications and recommendations

Previous reports have highlighted a possible relationship between groundwater levels in piezometer D1 and movements in inclinometer C1. Groundwater levels in piezometer D1 have previously shown a strong relationship with rainfall and this relationship should be specifically reviewed in future reports when data is available to refine understanding of that relationship.

The publicly accessible location of the acoustic inclinometer means it has become a target for vandalism, with youths seen by residents kicking the equipment. This has resulted in spurious data being recorded. It is recommended that any future acoustic installations are placed well-away from public accessible land to mitigate the risk of vandalism.

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Appendix A Digital data

